
Doctoral Dissertations

Student Theses and Dissertations

1972

Diffusiophoretic forces between juxtaposed atmospheric particles

Allen L. Williams

Follow this and additional works at: https://scholarsmine.mst.edu/doctoral_dissertations



Part of the [Physics Commons](#)

Department: Physics

Recommended Citation

Williams, Allen L., "Diffusiophoretic forces between juxtaposed atmospheric particles" (1972). *Doctoral Dissertations*. 1856.

https://scholarsmine.mst.edu/doctoral_dissertations/1856

This thesis is brought to you by Scholars' Mine, a service of the Missouri S&T Library and Learning Resources. This work is protected by U. S. Copyright Law. Unauthorized use including reproduction for redistribution requires the permission of the copyright holder. For more information, please contact scholarsmine@mst.edu.

DIFFUSIOPHORETIC FORCES BETWEEN
JUXTAPOSED ATMOSPHERIC PARTICLES

by

ALLEN LEE WILLIAMS, 1947-

A DISSERTATION

Presented to the Faculty of the Graduate School of the

UNIVERSITY OF MISSOURI - ROLLA

In Partial Fulfillment of the Requirements for the Degree

DOCTOR OF PHILOSOPHY

in

PHYSICS

1972

John C. Currence
Advisor

Samuel P. Leverston

[Signature]

Daryl Aloft

J. F. Stappert

J. M. Rivas

ABSTRACT OF THESIS

The role of diffusiophoretics and Stefan flow in droplet growth by accretion and in-cloud scavenging is examined. A calculation is made for the diffusive forces between two juxtaposed spheres one or both of which are undergoing diffusional growth and the influence of the diffusive force on the trajectories of two spheres, assumed to obey the creeping flow equations is found. The results are: a) the diffusiophoretic force between two water droplets is zero. b) the diffusiophoretic force between a spherical ice particle and a water droplet, although non-zero, is too small to appreciably influence the fall trajectories. c) the diffusiophoretic force between a growing, spherical cloud element and a spherical inert particle is large enough to enhance the collision efficiencies of the spheres.

ACKNOWLEDGEMENTS

The author wishes to thank Dr. J. C. Carstens for his advice and encouragement throughout the course of this study. Thanks are also due Dr. J. L. Kassner, Jr., Director of the Graduate Center for Cloud Physics Research, the staff, and graduate students of the Center for their many helpful suggestions. The author also wishes to thank Kathryn Berkbighler for her assistance in computer programming.

This research was supported by the Atmospheric Sciences Section, Office of Naval Research, under THEMIS Grant N00014-68-A-0497.

TABLE OF CONTENTS

	Page
ABSTRACT.....	ii
ACKNOWLEDGEMENT.....	iii
LIST OF FIGURES.....	v
LIST OF TABLES.....	vi
I. INTRODUCTION.....	1
II. REVIEW OF LITERATURE.....	6
A. Diffusiophoresis.....	8
B. Hydrodynamic Results.....	14
III. THEORY.....	17
A. Temperature and Vapor Density Profiles.....	24
B. Diffusiophoretic Force.....	30
IV. DISCUSSION OF RESULTS.....	36
A. Temperature and Vapor Density Profiles.....	36
B. Diffusiophoretic Force.....	41
C. Fall Trajectories.....	50
V. APPLICATIONS AND SUGGESTIONS FOR FURTHER STUDY.....	60
A. Accretion and In-cloud Scavenging.....	60
B. Limitations of Bispherical Geometry.....	62
C. Scavenging of Sub-micron Particles.....	64
VI. CONCLUSION.....	66
VII. BIBLIOGRAPHY.....	67
VIII. VITA.....	71
IX. APPENDIX.....	72
A. Computer Program.....	72

LIST OF FIGURES

Figures	Page
1. The Bispherical Coordinate System.....	25
2. Temperature Profile for Juxtaposed Growing Water Droplets..	37
3. Vapor Density Profile for Juxtaposed Growing Water Droplets.....	38
4. Temperature Profile for a Juxtaposed Water Droplet (Right) and an Ice Particle.....	40
5. Vapor Density Profile for a Juxtaposed Water Droplet (Right) and an Ice Particle.....	42
6. Temperature Profile for a Juxtaposed Water Droplet (Right) and an Inert Particle.....	43
7. Vapor Density Profile for a Juxtaposed Water Droplet (Right) and an Inert Particle.....	44
8. Fall Trajectories for 8 and 10 Micron Spheres.....	56
9. Fall Trajectories for 9.5 and 10 Micron Spheres.....	58

LIST OF TABLES

Table	Page
I. Diffusiophoretic Force Between an Ice Particle and a Water Droplet.....	46
II. Diffusiophoretic Force Between an Ice Particle and an Inert Particle.....	48
III. Diffusiophoretic Force Between a Water Droplet and an Inert Particle.....	49
IV. Values of Orientation and Instantaneous Velocity for Two Inert Spheres.....	51
V. Values of Orientation and Instantaneous Velocity for a 10 Micron Ice Particle (Sphere 1) falling toward an 8 Micron Water Droplet.....	54
VI. Values of Orientation for a 10 Micron Ice Particle (Sphere 1) falling toward an 8 Micron Inert Particle.....	55

I INTRODUCTION

The present study concerns micro-physical processes occurring near cloud elements that are several microns in radius. Such a cloud element, say a growing droplet, serves as a sink of water vapor causing vapor to diffuse toward the surface. Another aerosol particle placed near the growing droplet may experience a momentum transfer due to the vapor flow toward the droplet, the result being a force on the aerosol particle. Although the calculation of this diffusive force is the main concern in this study, other related transport phenomena such as heat transfer due to the heat released by vapor condensation and the momentum transfer problem resulting from the gravitational settling of the particles will be considered as well.

One incentive for pursuing such an undertaking is to understand the processes leading to coagulation of atmospheric aerosols. The study of the coagulation processes is commonly divided into the study of the motion leading to collision and the determination of the chances of coagulation after collision. Situations contemplated usually involve only two bodies. The present study will be confined to the motion of particles without consideration of the actual collision process.

The results of calculations to determine the trajectories of two spheres moving in a viscous fluid under the assumptions of steady, creeping motion have shown that the spheres have a very slight chance of colliding. According to an analysis by Davis¹ of the resistance on each of two spheres approaching each other along the line joining

centers, the resistance increases from the Stokes value for an isolated sphere at separation between surfaces of about 10 radii to a value of about 60 times the Stokes resistance at a separation of .01 radii. For this reason two spheres that are on an apparent collision course may approach to a very small separation between surfaces, then "slide" around each other without colliding.

Of a particular concern in cloud physics is the development of non-freezing clouds. Under typical cloud conditions a droplet, once activated, can readily grow by diffusion to several microns radius, after which the radial growth becomes very slow. The above hydrodynamic considerations suggest that the growth of droplets by accretion is very sluggish for these small droplets. It is therefore not apparent how the cloud droplet size spectrum develops from droplets with radii of a few microns to sizes large enough to grow effectively by accretion.

Another atmospheric process that appears to be retarded by the nature of the hydrodynamic forces between two falling spheres (at least in the size range from a few microns radius to around 25 microns radius) is the scavenging of atmospheric particulates by cloud elements, called rainout for in-cloud scavenging and washout for below-cloud scavenging. Since the viscosity of water is much greater than that of air, the fluid streamlines near a water droplet are the same as those around a solid sphere of the same radius. (For droplets of the sizes considered, there is no appreciable internal circulation). The above hydrodynamic considerations, then, suggest that the collisions between either droplets or (spherical)

ice particles and (spherical) atmospheric particles are highly improbable.

Since the spheres can approach to very close separations without colliding, any mechanism that alters the trajectories only slightly for small separations could greatly enhance the possibility of collision. This fact has led to identification of a few of these mechanisms. For larger spheres (approximately 25 microns in radius) their inertia causes them to cross the creeping flow streamlines thus, in some cases, effecting collision. Although the particle inertia changes the flow trajectory only slightly, the effect is most pronounced for small sphere separations. Another phenomenon that can enhance the number of collisions is electrical forces between particles. Again, the force attains its largest value for small sphere separations, although for multi-pole interactions there may be angular dependence as well.

If one or both of the spheres is a cloud element undergoing diffusional growth or evaporation, diffusive forces may influence the trajectories. Diffusive forces resulting from a gradient of concentration of at least one component in a multi-component gaseous system can be discussed in terms of Stefan flow and diffusiophoresis. Consider two parallel plates maintained at equal temperatures and separated by a humid air mixture. Let one of the plates serve as a source of water vapor and the other a sink, both being of equal strength. After steady-state is attained there will exist a non-zero average molecular velocity of the vapor toward the condensing plate, while the average molecular velocity of the air will be zero.

From the non-vanishing average molecular velocity of the vapor one can calculate a total mass velocity representing a flow of the medium toward the condensing surface. The average mass velocity of the medium is called the Stefan velocity named after its discoverer J. Stefan.²

Diffusive forces can also exist in the absence of a source or sink. Let the plates be replaced by vessels of equal temperature and pressure, each containing a different gas and connected by a tube of small diameter compared to the size of the vessels. If this system is allowed to attain steady-state, the total pressure in the tube will be constant and there will exist equal and opposite concentration gradients of the two gases. If the molecular weights of the two gases are not equal there will be a mass flow in the direction of the heavier gas component. A small particle placed in this tube would experience a force due to the gas flow around it. The term diffusiophoresis will be used to designate both this phenomenon and the similar situation resulting from Stefan flow.

The diffusiophoretic force on a particle depends on the value of the vapor density gradient established in the neighborhood of the particle. For an isolated growing cloud element the density gradient has its largest magnitude at the surface and falls off radially. If a particle is placed near the cloud element, it should experience a diffusive force due to the flow of water vapor toward the growing cloud element, the force being larger for smaller separations. (Neglected here is the possibility that the presence of the particle may change the vapor density profile.) To the extent that the

diffusive forces oppose the hydrodynamic forces caused by gravitational settling, the fall trajectories of the particles will be influenced in favor of collision.

The present effort is directed toward determining the extent to which diffusiophoretic forces affect the two-sphere trajectories from the standpoint of continuum hydrodynamics. An attempt will be made to keep the analysis general in the sense that a juxtaposition of two particles both undergoing diffusional growth or evaporation (denoted as volatile particles) will be considered, as well as the situation in which one of the particles is volatile the other inert (not undergoing diffusional growth or evaporation). The parameters that identify the particles such as the thermal conductivities and latent heats, will be kept distinct so that, for example, if both spheres are volatile one may correspond to a water droplet, the other an ice particle.

II REVIEW OF LITERATURE

The calculations of diffusiophoretic forces differ depending upon the particular Kundsens number regime considered.³ The Kundsens number is defined as $K = \lambda/R$ where R is the radius of the sphere suspended in a fluid of mean-free path λ . In the regime of large Kundsens number, the free-molecule regime, where $K \rightarrow \infty$ the transport processes occurring in the system can be described using uniform kinetic theory. This type of analysis can be extended to the transition regime ($10 > K_n > 10^{-1}$) by using the kinetic theory of non-uniform gases. In this regime, the distribution of molecules in the medium is influenced by the presence of the particle surface. This problem is further complicated by the lack of a reliable theory for the distribution of molecules leaving the surface.

In the continuum regime, defined in the limit as K_n approaches zero, the transport phenomena in the fluid can be described by the laws of continuum mechanics. It has been found by comparison with experiment that the continuum approach can be extended to the so called slip flow regime ($10^{-1} > K_n > 10^{-2}$) by using velocity slip and temperature jump boundary conditions.

If there exists a density gradient of one component of a binary mixture tangent to any surface in the system, the possibility of diffusion slip along this surface must be considered. Diffusion slip was first investigated by Kramers and Kistemaker⁴ (1943). They measured the slip velocity tangent to a wall along which was established a density gradient. Using a kinetic theory calculation, they found an expression for the diffusion slip coefficient. In the

usual continuum analysis of a gas flowing tangential to a stationery surface, the gas velocity evaluated at the surface is taken to be zero. However, the velocity of a gas diffusing along a surface is not always taken as zero. Diffusion slip represents a compromise in that, if a density gradient exists tangential to a surface, the velocity of the gas at the surface will take on some value between zero and the Fick's Law velocity. This diffusion slip phenomena influences the way in which momentum is transferred to a particle present in a diffusing gas mixture and thus may be of importance for the description of the diffusiophoretic force on a particle in the slip-flow and continuum regimes.

A. DIFFUSIOPHORESIS

Early research into what is now called diffusiophoresis was reported by Aitken⁵ (1883) whose results were later confirmed by Watson⁶ (1936). They essentially observed a dust free zone near an evaporating surface. Facy^{7,8,9} (1955) observed not only a dust free space near an evaporating droplet, but an increase in the number of dust particles near a condensing surface and a general movement of the particles toward the surface, with some being absorbed.

Deryagin and Dukhin¹⁰ (1956) first calculated the force on an inert particle in a constant density gradient. They considered the velocity field of the medium to be the Stefan velocity which, in this case, is the average mass velocity of the gaseous medium. They found the velocity of the particle to be the same as the Stefan velocity far removed from the particle. Deryagin and Dukhin¹¹ also calculated the force on a volatile particle in a diffusing gas mixture. In this case they found that in addition to the Stefan flow force, there is another force arising from the density gradients near the particle, causing a force opposing the Stefan flow. They concluded that, if the external concentration gradient is set up by a phase change of the same material as the droplet, the total force on the droplet is zero. (This result was later contradicted in a calculation by Deryagin and Yalamov.²⁰) The force on a volatile particle in both an externally imposed temperature and density gradient was found by Dukhin and Deryagin.¹² The velocity field of the gas medium was found by solving the Navier-Stokes equation for

creeping, steady motion, the equation of continuity for incompressible flow, and the steady-state, non-convective forms of Fick's law of diffusion, and Fouriers law of heat conduction simultaneously, subjected to the usual continuum boundary conditions (neglecting diffusion slip). According to this analysis the force on two widely separated water droplets is zero.

Waldmann (1961)¹³ found expressions for the forces on an inert particle in a diffusing gas mixture. Diffusion slip was accounted for in this analysis. The force on a particle in a binary gas mixture, each component of which is diffusing in opposite directions, was found to be in the direction of the heavier gas component. He also examined the situation of an inert particle suspended in a carrier gas through which vapor is diffusing. Although these results gave the proper direction for the forces, they do not agree with the experimental results of Schmitt and Waldmann^{14,15} and of Schmitt.¹⁶ Waldmann, therefore, proposed an empirical expression for the diffusion slip factor which led to reasonable experimental agreement.

Brock¹⁷ (1963) derived a new expression for the diffusion slip coefficient and using it calculated the force on an inert sphere suspended in a diffusing gas mixture. This analysis gave results very similar to those presented by Waldmann and can be shown to give identical results if the diffusion slip coefficients are set equal. Brock's slip coefficient depends on the accommodation coefficients of each species to the particle surface. Since there is no experimental information available on the accommodation coefficients of each chemical species in a mixture, the question of

diffusion slip is still open.

A completely different approach to the problem of diffusio-phoresis has been taken by Deryagin et. al.^{18,19,20} (1966). In this approach they considered two vessels containing binary gas mixtures, between which is maintained concentration and pressure differences. The vessels are separated by an "aerosol partition". Using irreversible thermodynamic arguments, they calculated the force on the aerosol particles in the partition. Diffusion slip was regarded as only of minor importance and was neglected in this development. The force calculated in this way gave reasonable agreement with the experimental value of the repulsive diffusio-phoretic force between a silver coated glass sphere and an evaporating droplet as reported by Prokhorov and Leonov.²¹

There are thus two general approaches to the problem of calculating diffusio-phoretic forces. The approach taken by Waldmann and Brock, as well as the earlier attempts by Deryagin's group, are based on hydrodynamic considerations, the basic differences being related to the question of diffusion slip. The non-equilibrium thermodynamic approach appears to be quite different in the method of attack. Unfortunately there is a scarcity of experimental data to determine which of the theories for calculation of the diffusio-phoretic force on large particles is correct. Even the question of the effect of diffusion slip on the particle motion has not been satisfactorily resolved. With regard to the irreversible thermodynamic approach, Waldmann²² has raised some questions regarding certain boundary conditions used.

For the problem at hand the conventional hydrodynamic approach will be exploited, admitting, however, that the irreversible thermodynamic approach may lead to equally suitable results. The existence of diffusion slip will be acknowledged and used in the subsequent development. By setting the slip coefficient equal to zero, the condition of no slip at the particle surface can be realized.

Corresponding to the diffusiophoretic forces which occur due to density gradients in the system, there exist thermophoretic forces which result from thermal gradients in the system. Due to the similarities of these two phenomena, thermophoresis has been developed by essentially the same researchers as diffusiophoresis. For the systems considered in this work, the thermophoretic forces are much smaller^{23,24} than the diffusiophoretic forces and will therefore be neglected in the present work.

There has been very little experimental work done to determine the continuum motion of particles in a diffusion field. Prokhorov and Leonov²¹ have presented results for the long-range diffusion forces between water droplets and non-volatile particles. They were, however, troubled by convection currents due to the cooling of the droplet from evaporation. To avoid the convection problem they had to heat the droplet to the same temperature as the surrounding medium. There has been no work reported for the diffusion forces on closely spaced spheres in the continuum regime.

Some work has been done to determine the atmospheric significance of diffusiophoresis for small particles. Goldsmith et. al.²⁵ found an experimental expression for the velocity of deposition of

small particles (about 1 micron radius) on a sulfuric acid surface. Based on this velocity, Goldsmith and May²⁶ presented a calculation to determine the role of diffusiophoresis in the scavenging of these particles from the atmosphere by growing water drops. They concluded that diffusiophoresis plays a very minor role in atmospheric scavenging. In this calculation they did not consider the drops to be falling due to gravity. Since a falling drop comes into proximity with more particles than a stationery drop and the diffusiophoretic force is larger for close separations, it is not clear that gravitational fall and the subsequent hydrodynamic problem can be ignored. Also neglected here is the heat liberated at the droplet surface due to condensation.

Using an analogy between the vapor density profile about a growing drop and the electric potential around a conducting sphere, Podzimek²⁷ investigated the binding of small AgI particles onto growing droplets and ice crystals. The distribution of vapor around the droplets and ice crystals was found by measuring the electrostatic field in an electrolytical vessel. The velocity of the particles was then calculated by using the Goldsmith velocity expression. The results indicate that the scavenging of small atmospheric particles may be important in mixed clouds. Again the results were not interpreted for a falling droplet or ice crystal.

According to the laboratory work of Vittori and Prodi²⁸, ice particles are 16 times more efficient than water droplets in scavenging small atmospheric particles. They also reported²⁹ that the observed velocity of particles in a density gradient is slower than

the Stefan flow.

Field experiments to examine the concentrations of different sized solid, water-insoluble particles in rainwater and hailstones were reported by Rosinski.³⁰ He concluded that "Stefan flow is probably the predominant mechanism of in-cloud scavenging of particles 1.5 - 5 microns in diameter when solid (ice crystal) and liquid (supercooled droplet) phases are present simultaneously."

B. HYDRODYNAMIC RESULTS

As a result of several contributions, the hydrodynamic problem of the steady, creeping motion of two spheres settling in a viscous fluid has been solved for different sized spheres and for arbitrary orientation. The assumptions of steady, creeping motion will be examined in the next section. The general motion of the spheres has been divided into motion parallel to the line of centers and motion perpendicular to the line of centers. Some remarks on this division of the velocity field will also be presented later.

An attempt to determine the motion of two spherical atmospheric particles moving in a viscous fluid under the assumptions of steady, creeping flow and in the absence of diffusion was reported by Hocking in his Ph.D. thesis³¹ (1958). Hocking's method of solution involved solving a large (in the limit infinite) set of simultaneous equations. The approximate solution was found by solving the first several members of the set of equations. The often quoted result of Hocking's work is that spheres with radii less than 19 microns do not collide.

An exact expression for the steady, creeping motion of two spheres moving with equal velocities parallel to their line of centers was derived by Stimpson and Jeffery³² (1926). Maude³³ (1961) and Brønner³⁴ (1961) independently extended the Stimpson and Jeffery solution to include the motion of two spheres moving with arbitrary velocities along their line of centers. Using this solution along with the solution used by Hocking for motion perpendicular to the line of centers (probably due to Wakiya³⁵), Davis and Sartor³⁶ (1967)

calculated collision efficiencies which were quite different from Hocking's original deduction. The new calculation failed to show the 19 micron cut-off. However, collision efficiencies for spheres of less than 20 microns in radius were still found to be quite small.

The mathematical analysis for a sphere moving parallel to a rigid infinite plane in steady, creeping flow was presented by O'Neill³⁷ (1964) based on a previous calculation by Dean and O'Neill³⁸ in 1963. The problem was done in bispherical coordinates enabling Wakiya³⁹ (1967) to generalize the problem to include two equal or unequal spheres moving perpendicular to their line of centers. Some numerical results for this motion were presented by Davis⁴⁰ (1969) who used the solution given by Wakiya.

A whole chapter in the book by Happel and Brenner⁴¹ (1965) is devoted to the discussion of the two sphere hydrodynamic problem. Calculations for the motion of two arbitrarily oriented equal spheres, as well as comparison with experimental results is given by Goldman, Cox, and Brenner⁴² (1966). The agreement with the experimental results of Eveson, Hall, and Ward⁴³ (1959) for which the Reynolds number never exceeded .018 is very good, to within 5 percent for values of a/h ranging from .097 to 1.0 where a is the sphere radius and h one-half the distance between centers. Theory and experiment both reveal that the small amount of rotation that does occur affects the sphere trajectories very slightly.

Other experiments on the settling of two equal, arbitrarily oriented spheres were done by Mathews and Smith⁴⁴, Slack and Mathews,⁴⁵ and by Jayaweera, Mason, and Slack.⁴⁶ These investigators observed

the upper sphere to overtake the lower sphere which does not agree with theoretical predictions for viscous flow around, equal spheres. The discrepancy is apparently inertial in origin. This is borne out by Klett⁴⁷ in his thesis where the Oseen correction is applied to two-sphere motion. Klett concludes that the inertial effects are important even for Reynolds numbers as low as .03 which corresponds to water droplets in air of radius 15 microns. Care must be taken when extending these experimental results, which were taken for solid spheres settling in a liquid medium, to atmospheric situations where the medium is gaseous.

III THEORY

The gas-particle system considered here will consist of spherical cloud elements of approximately 10 microns in radius suspended in a humid air environment. Since the molecular mean free path of the gas is about 6×10^{-6} cm., the Knudsen number is much smaller than unity. In view of this, the transport properties of the system can be described by the laws of continuum mechanics. Except for diffusion slip, the usual continuum boundary conditions will be employed. Although diffusion slip will be included, the complete continuum approach can be realized by setting the slip coefficient equal to zero. Due to the uncertainty regarding the continuum analysis of the fluid flow between two closely separated spheres, separations of less than about one micron will not be considered.

In order to find the force on particles suspended in a fluid, the mass velocity, \bar{v} , and the pressure, P , in the fluid must be established. The stress tensor can then be calculated from these fields. By integrating the stress tensor over the surface of the particles in the fluid, the force on the particles can be found.

Consider two spherical particles, one or both of which are volatile, suspended in an incompressible fluid, the whole system being in a constant gravitational field. The equation of motion of the fluid is:

$$\rho \frac{\partial \bar{v}}{\partial t} + \rho (\bar{v} \cdot \nabla) \bar{v} = - \nabla P + \eta \nabla^2 \bar{v} + \rho \bar{g} \quad , \quad (1)$$

where ρ is the total mass density of the fluid, η the fluid viscosity, \bar{g} the acceleration of gravity and t the time. The

equation of continuity for incompressible flow implies

$$\bar{\nabla} \cdot \bar{v} = 0. \quad (2)$$

The equation for heat transfer can be written⁴⁸

$$\frac{\partial T}{\partial t} + (\bar{v} \cdot \bar{\nabla}) T = \chi \nabla^2 T. \quad (3)$$

Here T is the temperature at any point in the system and $\chi = K/(\rho C_p)$, where K is the thermal conductivity at the position considered with mass density ρ and specific heat at constant pressure C_p . Diffusion of the water vapor of mass density ρ_v can be described by the equation:⁴⁹

$$\frac{\partial \rho_v}{\partial t} + (\bar{v} \cdot \bar{\nabla}) \rho_v = D \nabla^2 \rho_v \quad (4)$$

where D is the diffusion coefficient. According to a perturbation analysis by Brock^{17,50,51} the second term in both equations (3) and (4) can be neglected for the present system.

If one defines characteristic parameters such as the radius of one of the spheres, R (for closely spaced spheres the characteristic length may be best described as the distance between surfaces), and the free stream velocity, V , far away from both spheres, the Navier-Stokes equation can be expressed in terms of these quantities. The ratio of the inertial term to the viscous term then becomes $Re = \rho V R / \eta$, where Re is the Reynolds number. For an isolated sphere of 10 microns radius falling in air, the Reynolds number is of order 10^{-2} . Using this value for the present situation, the inertial term in the equation can be neglected in favor of the viscous term. Further justification for dropping the inertial term

appeals to the accuracy of such a calculation for falling, equal spheres when compared with experiment as discussed earlier.

Since the second term on the left of equation (3) can be neglected the heat equation becomes:

$$\frac{\partial T}{\partial t} = \chi \nabla^2 T.$$

A solution to this equation for a point source is given by Landeau and Lifshitz⁴⁸. The characteristic time required for heat to be conducted to a distance ℓ from the source is $\tau \sim \ell^2 / \chi$. This time can also be interpreted as the time necessary for the temperature of a non-uniformly heated body of dimension ℓ to become more or less constant throughout. Calling ℓ the distance between surfaces of two juxtaposed spheres, the temperature profile between the spheres reaches a steady value in a time $\tau = \ell^2 / \chi \sim 10^{-4}$ sec for $\ell = 50$ microns. During this time the spheres will settle to a new position resulting in a slight change in the temperature profile. This change in the profile with respect to time is very small compared to the changes contemplated for the whole trajectory. Therefore, the usual steady-state approximation will be made yielding for the heat conduction equation:

$$\nabla^2 T = 0. \quad (3')$$

The diffusion equation has exactly the same form and the above analysis applies to it as well. The mass diffusion equation therefore reduces to:

$$\nabla^2 \rho_v = 0. \quad (4')$$

Taking the curl of each term in the Navier-Stokes equation

(equation 1) and defining the vorticity as:

$\Omega = \bar{\nabla} \times \bar{v}$, equation 1 reduces to:

$\frac{\partial \Omega}{\partial t} = \eta / \rho \nabla^2 \Omega$. This has the same form as the heat conduction equation, so the relaxation time needed for the vorticity to reach its steady-state value is $\tau \sim \ell^2 \rho / \eta$. For a given value of ℓ this relaxation time is of the same order of magnitude as for heat conduction and diffusion, so the same argument for the steady-state approximation can be used. (Note that as the sphere separation is decreased the approximation becomes stronger.) A particle moving in steady flow is necessarily moving in such a way as to make the total force on the particle zero. Thus for the situation considered here, the gravitational force on the particle and the hydrodynamic forces must sum to zero.

Since the density of the particles is much greater than that of their environment the buoyancy force on the particles, represented by the last term in (1), can be neglected. The equation of motion of the fluid therefore reduces to:

$$\bar{\nabla} P = \eta \nabla^2 \bar{v}. \quad (1')$$

As pointed out by Berker⁵² and Brenner,⁵³ since these equations of motion and the boundary conditions (to be discussed later) are linear functions of velocity and pressure, the problem of the general motion of two spheres satisfying equations (1') and (2) can be decomposed into a number of simpler problems. For example, the velocity can be written as the sum of other velocities say $\bar{v} = \bar{v}_1 + \bar{v}_2$ and the pressures corresponding to each of these velocities become $P = P_1 + P_2$. If these parameters satisfy the

equations

$$\overline{\nabla P}_1 = \eta \nabla^2 \bar{v}_1, \quad \overline{\nabla P}_2 = \eta \nabla^2 \bar{v}_2,$$

$\nabla \cdot \bar{v}_1 = 0$, and $\nabla \cdot \bar{v}_2 = 0$ subject to boundary conditions consistent with the boundary conditions of the total flow field, then solving each of the sub-problems is equivalent to solving the original problem. Brenner also shows that a stress tensor and hence a force, can be found from each of the individual flow fields, the total force being a sum of each of the smaller forces. For the problem at hand, it is convenient to separate the hydrodynamic problem into three parts. These are motion perpendicular and parallel to the line of centers due to gravitational settling, and motion due to diffusiophoretic forces. The problem of gravitational settling has been solved, as discussed in the literature review. Once the diffusiophoretic force is known, the velocity of each particle for any separation can be found by requiring that the total force on the particles be zero.

Consider then two juxtaposed particles, at least one of which is growing by diffusion, positioned in an otherwise quiescent fluid. From the axisymmetric geometry of this configuration, any force present must be along the line of centers of the spheres. The spheres will be constrained not to move and the necessary force of constraint will be found by solving equations (1') and (2) for the velocity and pressure fields due to diffusion. At the surface of a volatile particle the boundary conditions on the velocity are:

$$\bar{v} \cdot \bar{n} \Big|_{\text{SURFACE}} = \frac{-D}{\rho_g} \left(\overline{\nabla \rho_v} \cdot \bar{n} \right) \Big|_{\text{SURFACE}} \quad (5)$$

and

$$\bar{v} \cdot \bar{\tau} \Big|_{\text{SURFACE}} = - \frac{\sigma D}{\rho_g} \left(\bar{\nabla} \rho_v \cdot \bar{\tau} \right) \Big|_{\text{SURFACE}} \quad (6)$$

Here \bar{n} and $\bar{\tau}$ are unit vectors normal and tangential to the particle surface respectively, ρ_v the mass density of the water vapor, D the diffusion coefficient of the binary gas system, ρ_g the density of the dry air, and σ the diffusion slip coefficient first derived by Kramers and Kistemaker.⁴ For an inert particle no vapor can penetrate the surface. Writing the condition of impenetrability as

$$\bar{\nabla} \rho_v \cdot \bar{n} \Big|_{\text{SURFACE}} = 0 \quad , \quad (7)$$

it is apparent that the fluid velocity normal to the particle surface as described by equation (5) must be zero. The tangential velocity will have the same form as shown in equation (6). In fact, the velocity boundary conditions for both inert and volatile particles can be described by equations (5) and (6), the difference being that the density profile will be different in each case.

Neglecting temperature jump at the particle surface, the temperature at the surface of both inert and volatile particles must be continuous, that is:

$$T_e \Big|_{\text{SURFACE}} = T_i \Big|_{\text{SURFACE}} \quad , \quad (8)$$

where T_e is the temperature external to the particle and T_i the internal temperature.

A droplet that is undergoing diffusional growth must conduct away any heat released at the surface either via the outside medium or into the droplet. This can be expressed in terms of a steady-state power balance:

$$\bar{n} \cdot \left\{ LD \bar{\nabla} \rho_v + K_e \bar{\nabla} T_e - K_i \bar{\nabla} T_i \right\} \Big|_{\text{SURFACE}} = 0. \quad (9)$$

Here L is the latent heat of the volatile particle, K_e the thermal conductivity external to the particle, and K_i that internal to it. If the particle is inert the condition of impenetrability reduces equation (9) to $\bar{n} \cdot \{ K_e \bar{\nabla} T_e - K_i \bar{\nabla} T_i \} = 0.$ (10)

It will be assumed that the temperature and vapor density are in equilibrium at the droplet surface. This assumption is good under the condition that $(\lambda/R \delta) \gg 1$ (see Carstens and Kassner,⁵⁴ 1968) where δ is the sticking coefficient. (Actually these same criteria determine the validity of the thermal continuity boundary condition as well.) The condition of thermal equilibrium, therefore,

$$\text{becomes: } \rho_v \Big|_{\text{SURFACE}} = bT \Big|_{\text{SURFACE}} + c \quad (11)$$

where b and c are constants determined from experiment.

Far from the particles, the temperature and vapor density must attain their bulk values so: $\rho = \rho_\infty$ and $T = T_\infty$ at an infinite distance from the spheres. Also at an infinite distance from the spheres, there can be no fluid motion caused by diffusion since there are no density gradients thus $\bar{v} = 0$ at infinity.

In the more general situation the differential equations describing the temperature, vapor density, velocity, and pressure fields are coupled such that all four equations must be solved simultaneously. In the system considered here, approximations have been made which decouple the pressure and velocity fields

from the temperature and density fields. In what follows, the temperature and density profiles will be found by simultaneously solving equations (3') and (4'). These solutions will be subjected to boundary conditions (7), (8), and (10) at the surface of an inert particle and boundary conditions (8), (9), and (11) if the particle is volatile. Once the vapor density profile for the system is known, the pressure and velocity fields can be found by simultaneously solving equations (1') and (2) subject to boundary conditions (5) and (6).

A. TEMPERATURE AND VAPOR DENSITY PROFILES

Due to the geometry of the system, a bispherical coordinate system will be used to find the profiles. A review of bispherical coordinates may be found in Happel and Brenner⁴¹ or Morse and Feshbach.⁵⁵ Figure 1 shows the coordinate grid. The full three dimensional system is obtained by rotating this system about the z - axis.

The solutions to equations (3') and (4') in bispherical coordinates are:

$$T_e = T_\infty + \sqrt{f} \sum_{n=0}^{\infty} [A_n \exp(-\mu_n) + B_n \exp(\mu_n)] P_n(w) \quad \text{and} \quad (12)$$

$$\rho_v = \rho_\infty + \sqrt{f} \sum_{n=0}^{\infty} [C_n \exp(-\mu_n) + D_n \exp(\mu_n)] P_n(w) \quad \text{for } -\mu_2 \leq \mu \leq \mu_1, \quad (13)$$

$$T_1' = \sqrt{f} \sum_{n=0}^{\infty} F_n' \exp(-\mu_n) P_n(w) \quad \text{for } \mu \geq \mu_1, \quad (14)$$

$$T_1'' = \sqrt{f} \sum_{n=0}^{\infty} F_n'' \exp(\mu_n) P_n(w) \quad \text{for } \mu \leq -\mu_2. \quad (15)$$

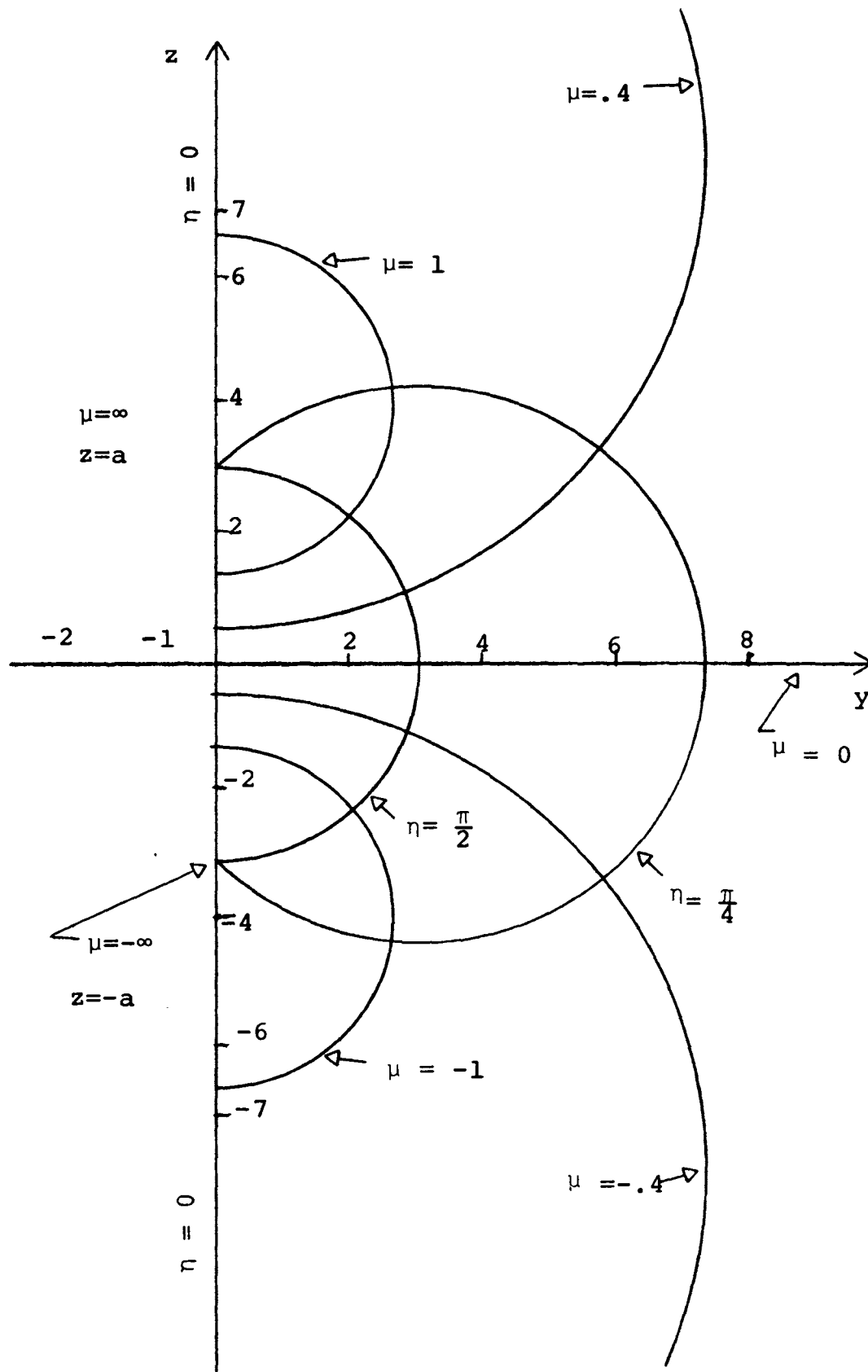


Figure 1. The Bispherical Coordinate System.

Here A_n , B_n , C_n , D_n , F_n' and F_n'' are arbitrary constants to be determined from the boundary conditions, ρ_∞ and T_∞ are the ambient values of temperature and vapor density, $\mu_n = (n + \frac{1}{2})$, $w = \cos \eta$, and $f = \cosh \mu - w$. μ_1 describes the surface of particle 1 with center on the positive z-axis and $-\mu_2$ describes the surface of particle 2 centered on the negative z-axis. (The minus sign associated with μ_2 will be carried along explicitly so μ_2 is a positive number.)

These equations have been solved and the profiles computed for two juxtaposed water droplets by Carstens, Williams, and Zung⁵⁶ (1970) with comments by Williams and Carstens⁵⁷ (1971). The calculations will be generalized here to include particles with distinct latent heats and thermal conductivities. Also, the profiles will be found for an inert and volatile particle juxtaposition.

For the situation in which both particles are volatile the boundary conditions yield:

$$A_n - F_n' + B_n \exp(2 \mu_{n1}) = -\sqrt{2} T_\infty \quad (16)$$

and

$$A_n \exp(2 \mu_{n2}) + B_n - F_n'' = -\sqrt{2} T_\infty \quad (17)$$

from thermal continuity at particles 1 and 2 respectively,

$$C_n - b_1 F_n' + D_n \exp(2 \mu_{n1}) = \sqrt{2} (C_1 - \rho_\infty) \quad (18)$$

and

$$C_n \exp(2 \mu_{n2}) + D_n - b_2 F_n'' = \sqrt{2} (C_2 - \rho_\infty) \quad (19)$$

from thermal equilibrium,

$$\alpha_n = n \alpha_{n-1}' + (n+1) \alpha_{n+1}' \quad (20)$$

and

$$\beta_n = n\beta'_{n-1} + (n+1)\beta'_{n+1} \quad (21)$$

from the power balance. The following notation has been used:

$\mu_{n1} = (n+\frac{1}{2})\mu_1$, $\mu_{n2} = (n+\frac{1}{2})\mu_2$, b_1 , b_2 , c_1 , c_2 are the constants obtained by requiring thermal equilibrium at each surface, K_{11} , K_{12} , K_e are the thermal conductivities of particle 1, particle 2, and the medium external to the particles, L_1 and L_2 , are the latent heats of condensation (sublimation) of particles 1 and 2, D the diffusion coefficient,

$$\alpha_n = T_{11}(C_n + \Gamma_{e1}A_n - \Gamma_{i1}F'_n)\exp(-\mu_{n1}) + T_{12}(D_n + \Gamma_{e1}B_n)\exp(\mu_{n1}),$$

$$\alpha'_n = -(C_n + \Gamma_{e1}A_n - \Gamma_{i1}F'_n)\exp(-\mu_{n1}) + (D_n + \Gamma_{e1}B_n)\exp(\mu_{n1}),$$

$$\beta_n = T_{22}(C_n + \Gamma_{e2}A_n)\exp(\mu_{n2}) + T_{21}(D_n + \Gamma_{e2}B_n + \Gamma_{i2}F''_n)\exp(-\mu_{n2}),$$

$$\beta'_n = (C_n + \Gamma_{e2}A_n)\exp(\mu_{n2}) - (D_n + \Gamma_{e2}B_n + \Gamma_{i2}F''_n)\exp(-\mu_{n2}),$$

$$T_{11} = \sinh(\mu_1) - (2n+1)\cosh(\mu_1),$$

$$T_{12} = \sinh(\mu_1) + (2n+1)\cosh(\mu_1),$$

$$T_{21} = \sinh(\mu_2) - (2n+1)\cosh(\mu_2),$$

$$T_{22} = \sinh(\mu_2) + (2n+1)\cosh(\mu_2),$$

$$\Gamma_{e1} = \frac{K_e}{L_1 D}, \quad \Gamma_{e2} = \frac{K_e}{L_2 D}, \quad \Gamma_{11} = \frac{K_{11}}{L_1 D}, \quad \Gamma_{12} = \frac{K_{12}}{L_2 D}.$$

Each of equations (16-21) yield a separate equation for each value of n . Consider the case for $n=0$, equations (16-19) give four equations and six unknowns. Equations (20) and (21) add two more equations and ~~six~~ more unknowns, since for $n=0$, α'_{n+1} and β'_{n+1} give the $n=1$ order constants. It appears that two more conditions are needed to specify the solution. This situation can be alleviated by summing the power balance boundary condition over every point on the surface of each of the spheres. For particle 1 this condition can be written:

$$\iint \bar{n} \cdot \{ \overline{\nabla \rho_v} \big|_{\mu=\mu_1} + \Gamma_{e1} \overline{\nabla T_e} \big|_{\mu=\mu_1} - \Gamma_{11} \overline{\nabla T'_1} \big|_{\mu=\mu_1} \} dS = 0$$

where \bar{n} is a unit vector normal to the surface. This condition yields at surface 1 and 2 respectively:

$$\sum_{n=0}^{\infty} (D_n + \Gamma_{e1} B_n) = 0, \quad (22)$$

$$\sum_{n=0}^{\infty} (C_n + \Gamma_{e2} A_n) = 0. \quad (23)$$

Consider the situation in which one of the particles is inert and the other volatile. Enforcing the volatile boundary conditions at surface 1 yields equations (16), (18), (20) and (22). At the surface of the inert particle thermal continuity holds, thus equation

(17) is valid. The condition of impenetrability, equation (7), enforced at surface 2 gives:

$$\gamma_n = n \gamma'_{n-1} + (n+1) \gamma'_{n+1} \quad (24)$$

where $\gamma_n = T_{22} C_n \exp(\mu_{n2}) + T_{21} D_n \exp(-\mu_{n2})$

and $\gamma'_n = C_n \exp(\mu_{n2}) + D_n \exp(-\mu_{n2})$.

The power balance condition at surface 2 can again be written as equation (21), however, some simplifications can be made due to the condition of impenetrability. Integrating the power balance and the condition of impenetrability over surface 2 yields:

$$\sum_{n=0}^{\infty} A_n = 0 \quad \text{and} \quad \sum_{n=0}^{\infty} C_n = 0 \quad . \quad (25)$$

These boundary condition equations are sufficient to specify unique values of the arbitrary constants in equations (12) - (15) for juxtapositions of both two volatile particles and one inert particle the other volatile. Numerical values for the constants were found on an IBM - 360 computer. The approximate temperature and vapor density profiles were then found by taking a sufficient number of terms in the summation expressions for the temperature and vapor density. The constants either remained the same or converged for increasing values of n , making possible a solution accurate to several decimal places by taking less than ten terms in the sums.

B. DIFFUSIOPHORETIC FORCE

Before calculating the force, it will prove expedient to simplify the velocity boundary conditions. This can be done by noting that the force calculation can be reduced to a number of sub-problems as discussed earlier. Consider the velocity division $\bar{v} = \bar{v}' + \overline{\nabla \phi}$ where ϕ is any scalar function. The pressure gradient corresponding to the velocity $\overline{\nabla \phi}$ and the resulting force are zero,¹³ so this amounts to a trivial division. It is then apparent that the gradient of any scalar function can be added to the velocity field without changing the force due to this velocity.

Consider the diffusional velocity field

$$\bar{v} = \nabla \frac{D\rho_v}{\rho_g} \sim \frac{D}{\rho_g} \overline{\nabla \rho_v}, \text{ where terms of order } \overline{\nabla \rho_v} \cdot \overline{\nabla \rho_g} / \rho_g^2 \text{ are}$$

neglected. This velocity can be added to the total diffusion velocity field changing the surface boundary conditions, equations (5) and (6), yielding:

$$\bar{v} \cdot \bar{n} \Big|_{\text{SURFACE}} = 0 \quad \text{and} \quad (26)$$

$$\bar{v} \cdot \bar{\tau} \Big|_{\text{SURFACE}} = (1 - \sigma) \frac{D}{\rho_g} (\overline{\nabla \rho_v} \cdot \bar{\tau}) \Big|_{\text{SURFACE}}. \quad (27)$$

Since the vapor density far from the particle surfaces is at its bulk value and, hence, constant the gradient velocity is zero. The condition on the velocity far from the surface, therefore, remains zero.

Since the diffusion flow field has axisymmetric geometry, a stream function representing the solution to equations (1') and (2) can be used. According to Stimpson and Jeffery³² the stream

function in bispherical coordinates for this situation can be written as:

$$\Psi = f^{-3/2} \sum_{n=0}^{\infty} U_n V_n(w) \quad (28)$$

$$\begin{aligned} \text{where } U_n = & A'_n \cosh(n - \frac{1}{2}) \mu + B'_n \sinh(n - \frac{1}{2}) \mu \\ & + C'_n \cosh(n + \frac{3}{2}) \mu + D'_n \sinh(n + \frac{3}{2}) \mu . \end{aligned}$$

Here A'_n , B'_n , C'_n , and D'_n are constants to be determined and

$$V_n(w) = C_{n+1}^{-\frac{1}{2}}(w) \quad \text{where } C_n^{-m} \text{ are the Gegenbauer Polynomials}$$

defined by:

$$\sum_{n=0}^{\infty} C_n^m(x) t^n = (1 - 2xt + t^2)^{-m}$$

The velocities normal and tangential to the sphere surfaces as related to the stream function are:⁴¹

$$\bar{v} \cdot \bar{n} = \frac{f^2}{a^2} \frac{\partial \Psi}{\partial w} \quad \text{and} \quad (29)$$

$$\bar{v} \cdot \bar{\tau} = \frac{f^2}{a^2 \sin \eta} \frac{\partial \Psi}{\partial \mu} . \quad (30)$$

Here a is a constant used in bispherical coordinates for a particular sphere separation defined as $a = (d^2 - r^2)^{\frac{1}{2}}$, where

d is the distance along the z-axis to the center of the sphere of radius r. The boundary condition for velocity normal to the surface in terms of the stream function becomes:

$$\left. \frac{\partial \Psi}{\partial w} \right|_{\text{SURFACE}} = 0 \quad \text{thus} \quad \left. \Psi \right|_{\text{SURFACE}} = \text{constant.} \quad \text{Since the stream}$$

function is zero on the axis of symmetry, which intersects the surface, this boundary condition on the stream function becomes:

$$\left. \Psi \right|_{\text{SURFACE}} = 0. \quad (31)$$

This can be written as:

$$\sum_{n=0}^{\infty} U_n V_n(w) \left|_{\text{SURFACE}} = 0 \quad \text{or} \quad \left. U_n \right|_{\text{SURFACE}} = 0,$$

since the Gegenbauer polynomials constitute a complete set and are independent for each n.

In terms of the stream function, boundary condition (27) becomes:

$$\left. \frac{f^2}{a^2 \sin \eta} \frac{\partial \Psi}{\partial \mu} \right|_{\text{SURFACE}} = (1 - \sigma) \frac{D}{\rho_g} \left(\bar{\nabla} \rho_v \cdot \bar{\tau} \right) \left|_{\text{SURFACE}}.$$

In bispherical coordinates the direction tangential to the surface is described by the unit vector in the η -direction thus:

$$\left. \frac{\partial \Psi}{\partial \mu} \right|_{\text{SURFACE}} = -K \frac{1-w^2}{f} \left. \frac{\partial \rho_v}{\partial w} \right|_{\text{SURFACE}}$$

where $K = \frac{(1-\sigma)Da}{\rho_g}$. From equation (13) the vapor density can

be written as:

$$\rho_v = \sqrt{f} \sum_{n=0}^{\infty} \phi_n P_n(w) \text{ so:}$$

$$\frac{\partial \rho_v}{\partial w} = \sqrt{f} \sum_{n=0}^{\infty} \phi_n P'_n(w) - \frac{1}{2\sqrt{f}} \sum_{n=0}^{\infty} \phi_n P_n(w) \text{ where}$$

$P'_n(w) = \partial P_n(w) / \partial w$. This boundary condition therefore becomes:

$$f^{3/2} \frac{\partial \psi}{\partial n} \bigg|_{\text{SURFACE}} = K(1-w^2) \left[\frac{1}{2} \sum_{n=0}^{\infty} \phi_n P_n(w) - \cosh \mu \sum_{n=0}^{\infty} \phi_n P_n(w) + w \sum_{n=0}^{\infty} \phi_n P'_n(w) \right] \bigg|_{\text{SURFACE}}$$

This equation can be expressed in terms of Gegenbauer polynomials by use of the identities:^{32, 58}

$$(1-w^2) P_n(w) = \frac{(n+1)(n+2)}{(2n+1)(2n+3)} V_{n+1}(w) - \frac{n(n-1)}{(2n+1)(2n-1)} V_{n-1}(w)$$

$$(1-w^2) P'_n(w) = \frac{n(n+1)}{2n+1} V_n(w), \text{ and}$$

$$(1-w^2)w P'_n(w) = \frac{n(n+1)(n-1)}{(2n-1)(2n+1)} V_{n-1}(w) + \frac{n(n+1)(n+2)}{(2n+1)(2n+3)} V_{n+1}(w).$$

By substituting these identities into the boundary condition equation and simplifying one obtains:

$$f^{3/2} \frac{\partial \psi}{\partial \mu} \bigg|_{\text{SURFACE}} = \sum_{n=0}^{\infty} \phi_n V_n(w) \bigg|_{\text{SURFACE}} \quad (32)$$

$$\text{where } \phi_n = \frac{K}{2} \frac{n(n+1)}{(2n+1)} \{ \phi_{n+1} + \phi_{n-1} - 2 \phi_n \cosh(\mu) \}.$$

Writing the expression for the stream function, equation

(28), as

$$f^{3/2} \psi = \sum_{n=0}^{\infty} U_n V_n(w),$$

taking the partial derivative with respect to μ of both sides of this equation, and evaluating the expression at the surface in light of equation (31) one obtains:

$$f^{3/2} \frac{\partial \psi}{\partial \mu} \Big|_{\text{SURFACE}} = \sum_{n=0}^{\infty} \frac{\partial}{\partial \mu} U_n V_n(w) \Big|_{\text{SURFACE}} .$$

Using this result the boundary condition for velocity tangential to the surface as expressed by equation (32) becomes

$$\sum_{n=0}^{\infty} \phi_n V_n(w) \Big|_{\text{SURFACE}} = \sum_{n=0}^{\infty} \frac{\partial U_n}{\partial \mu} V_n(w) \Big|_{\text{SURFACE}} \quad \text{or}$$

$$\phi_n - \frac{\partial U_n}{\partial \mu} \Big|_{\text{SURFACE}} = 0 . \quad (33)$$

Evaluating equations (31) and (33) at each particle surface the constants A'_n , B'_n , C'_n and D'_n can each be found for every n . As shown by Stimpson and Jeffery,³² for cases in which the stream function (28) is a solution to the equations of motion, the force on each sphere becomes:

$$F_1 = \eta \pi 2 \sqrt{2} / a \sum_{n=1}^{\infty} (2n+1) (A'_n + B'_n + C'_n + D'_n) \quad (34)$$

for the sphere $\mu = \mu_1$ and

$$F_2 = - \eta \pi 2 \sqrt{2} / a \sum_{n=1}^{\infty} (2n+1) (A'_n - B'_n + C'_n - D'_n) \quad (35)$$

for the sphere at $\mu = -\mu_2$.

By evaluating equations (31) and (33) at each particle surface, four independent equations are found from which the four constants

A'_n , B'_n , C'_n , and D_n for each value of n can be determined. Since the constants converge for large values of n , the force can be found to the desired accuracy by terminating the sums in equations (34) and (35) at an appropriate value of n .

The resistance force for two spheres moving parallel to their line of centers³² and moving perpendicular to their line of centers⁴⁰ is found in each case by assuming that the spheres have a given velocity, then calculating the resistance force resulting from this velocity. In the present situation, these results will be used to give the velocities both parallel and perpendicular to the line of centers that are necessary to make the total force in each of these directions zero. For motion parallel to the line of centers the sum of the diffusiophoretic force, the parallel component of the gravitational force, and the resistance force on each particle must be zero. For the motion perpendicular to the line of centers the perpendicular component of the gravitational force and the resistance force on each particle must add up to zero.

For a given two particle system at an arbitrary separation the velocity of each particle can be found from the above analysis. To find the trajectories the particles will be allowed to move at these velocities for an increment of time resulting in a new orientation for which new particle velocities can be found.

IV DISCUSSION OF RESULTS

A. TEMPERATURE AND VAPOR DENSITY PROFILES

The temperature and vapor density profiles for two juxtaposed, growing water droplets of equal radii are shown in figures 2 and 3. Corresponding profiles for unequal droplets were given by Carstens et. al.⁵⁶ In both cases the internal droplet temperatures are constant. Since the temperature and vapor density evaluated at the particle surface are related by the condition of thermal equilibrium, it follows that if the droplet temperature is constant the vapor density gradient tangential to the surface and evaluated at the surface is zero.

Some insight into the question of why the internal temperatures of juxtaposed droplets are constant can be found by considering an isolated growing droplet. The internal temperature for an isolated droplet is constant. If external temperature and density gradients are established near the isolated droplet, the internal droplet temperature will remain constant only if the imposed temperature and density gradients satisfy the power balance:

$$\{LD\bar{\nabla}\rho + K_e \bar{\nabla}T\} \cdot \bar{n} = 0. \quad (36)$$

Thus the total amount of heat released at the surface as a consequence of the externally imposed vapor flow must be carried away by the externally imposed temperature gradient. The temperature and density gradients outside any isolated droplet satisfy equation (36). Thus, if two droplets are juxtaposed and the isolated profiles of each serve as the externally imposed gradients on the other, it

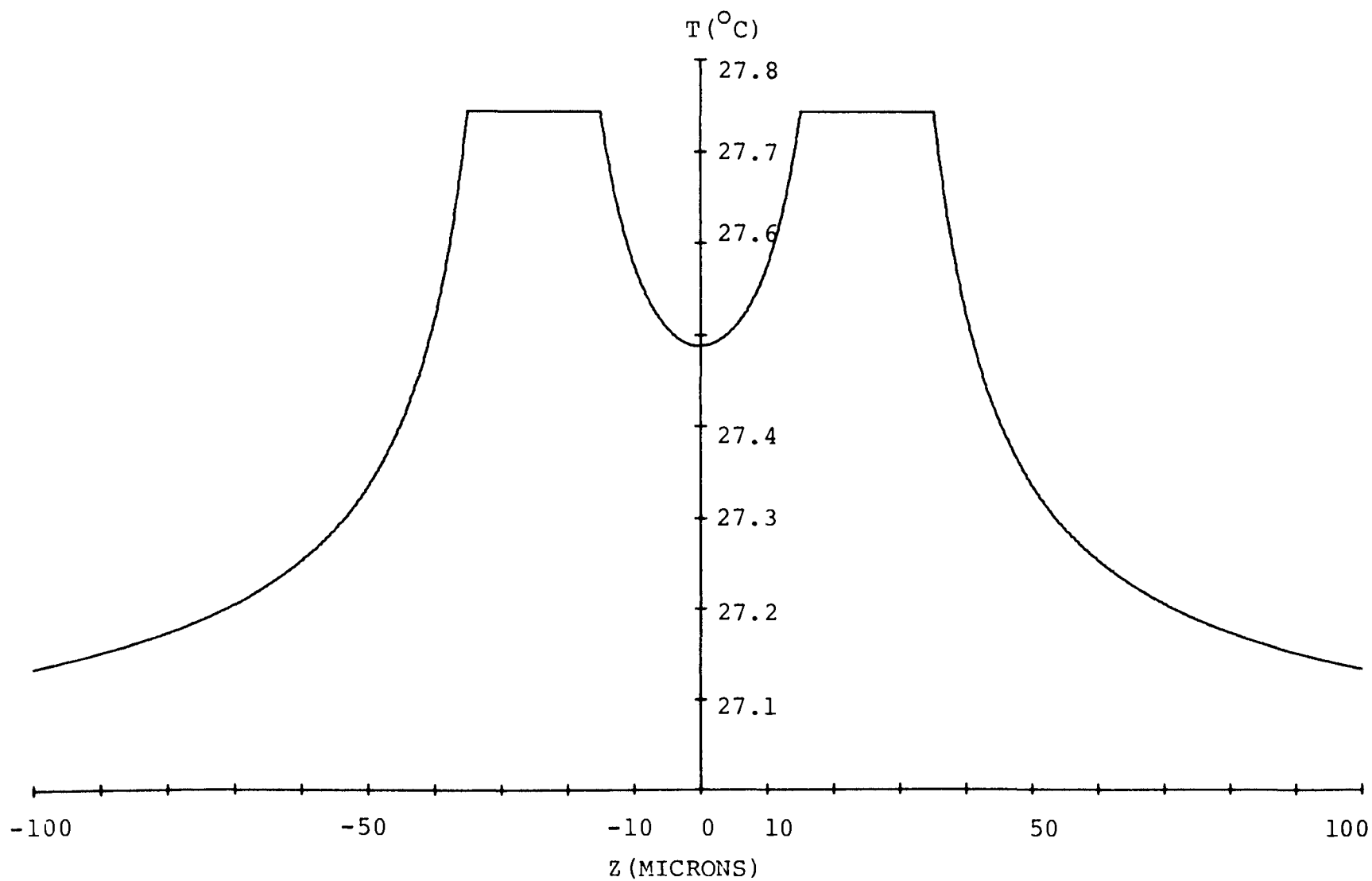


Figure 2. Temperature profile for juxtaposed growing water droplets. Supersaturation is $S \approx 1.05$.

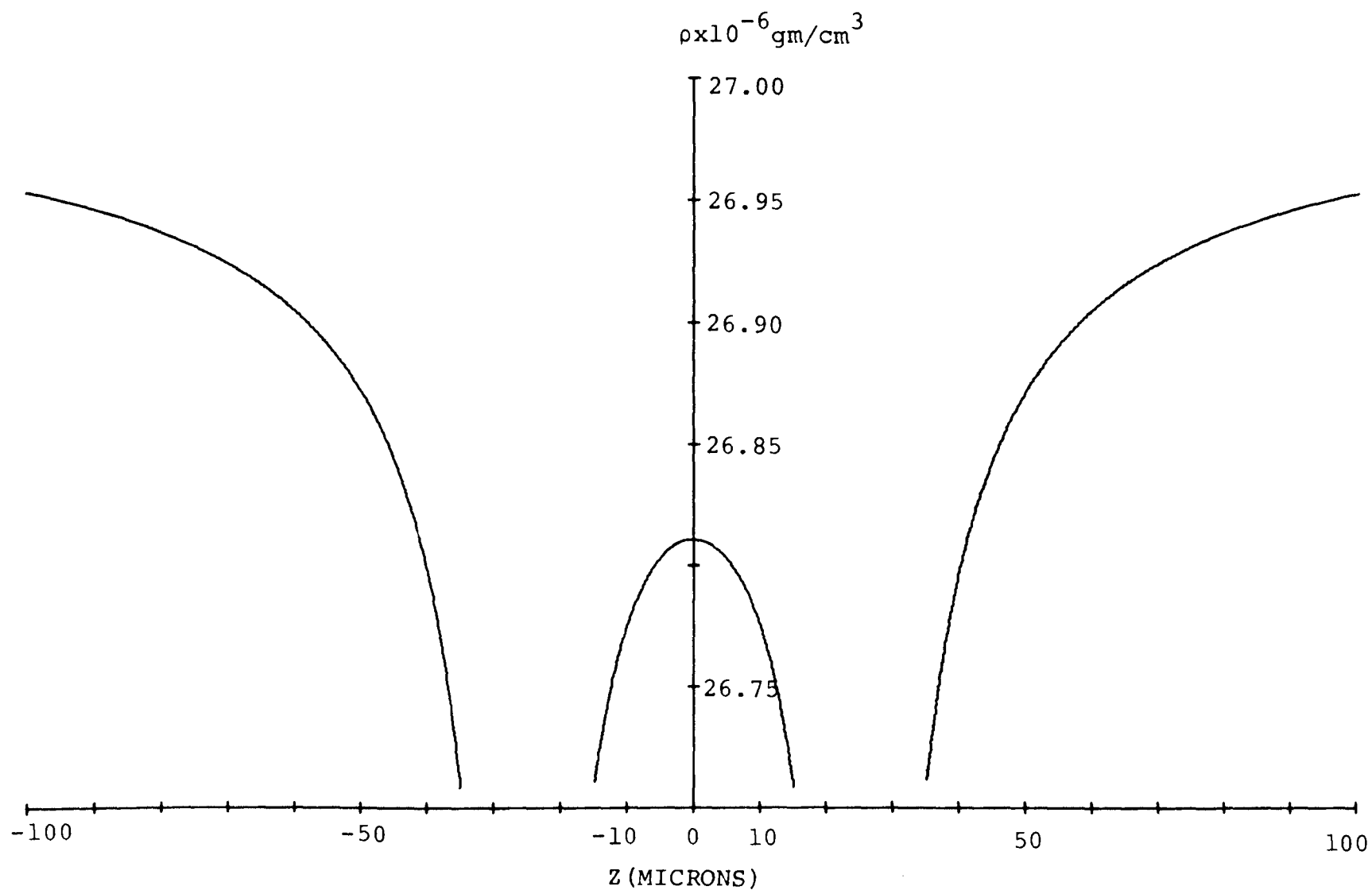


Figure 3. Vapor density profile for juxtaposed growing water droplets. $S = 1.05$.

is reasonable to expect the internal temperatures to remain constant. Apparently this kind of balance exists in the actual two droplet situation.

Using a similar analysis, consider the juxtaposition of an ice particle and a water droplet, both undergoing diffusional growth. The temperature and vapor density gradients outside an isolated ice particle should satisfy equation (36), where L in this case represents the heat of sublimation. Therefore, at any point outside the isolated ice particle there is a flow of vapor which, if allowed to deposit on an ice surface, would give up an amount of heat that would be totally conducted away by the external temperature gradient existing at that point. However, if a water droplet is placed near the ice particle, the side of the droplet farthest from the ice surface would intercept a vapor flow from the externally imposed density gradient, which upon condensation would give up a smaller amount of heat than is conducted away by the externally imposed temperature gradient, since the latent heat of water is less than the heat of sublimation of ice. This side of the droplet would thus become cooler. The other side of the droplet would be heated since the imposed temperature gradient conducts more heat toward the surface than is lost due to evaporation induced by the externally imposed density gradient. Figure 4 shows the temperature profile for a juxtaposed ice particle and a water droplet. Note that there is a gradient of temperature inside the water droplet and in the sense as suggested. The above analysis indicates

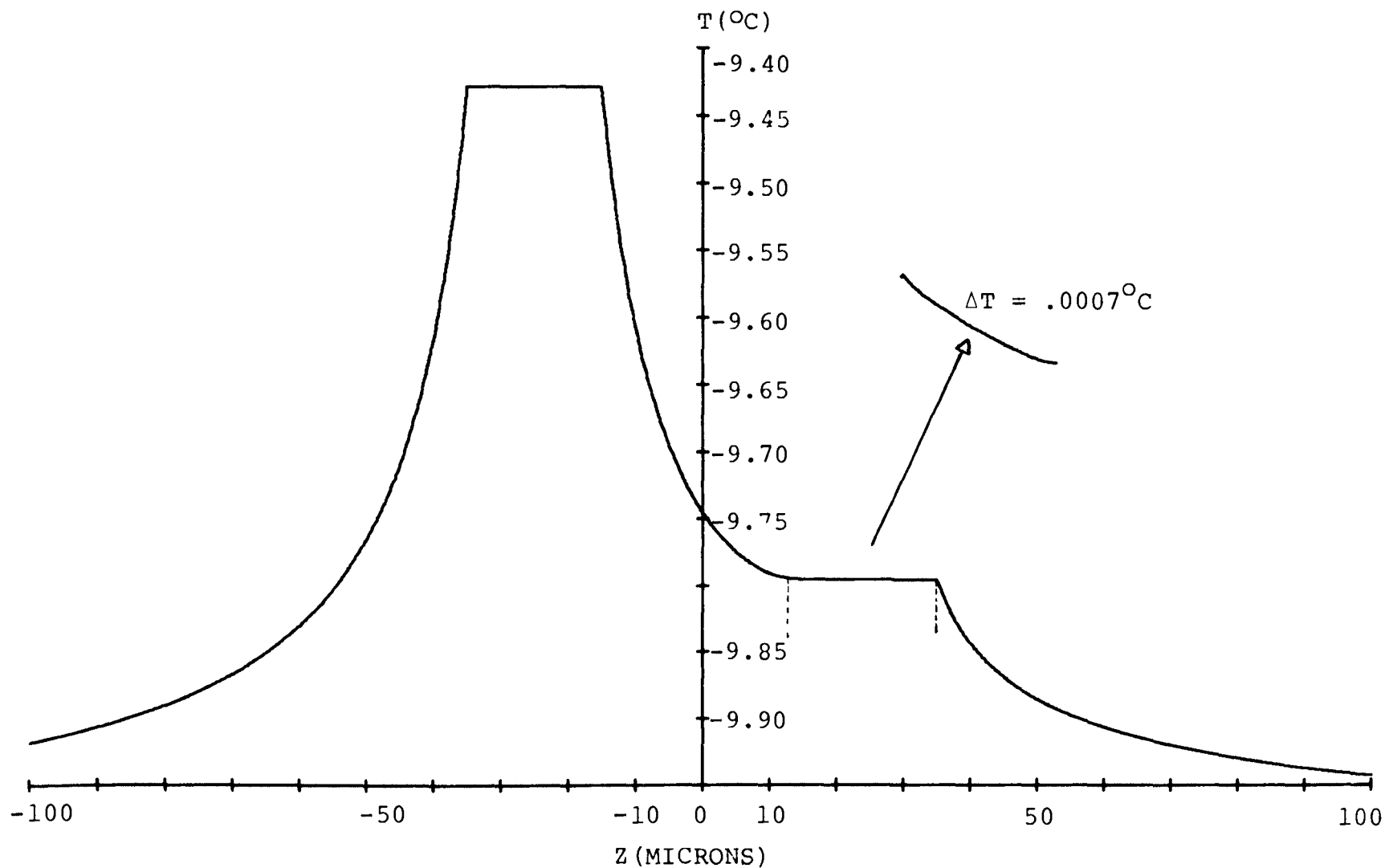


Figure 4. Temperature profile for a juxtaposed water droplet (right) and an ice particle. The supersaturation with respect to the water is 1.05 and with respect to the ice 1.137. The insert shows the internal temperature of the water.

that there should also exist a temperature gradient internal to the ice particle which turns out to be true. However, this gradient is much smaller, partly due to the higher thermal conductivity of ice. Figure 5 shows the corresponding vapor density profile for the ice water juxtaposition.

The temperature and vapor density for a system of ice and an inert particle is shown in figures 6 and 7 respectively. It is apparent from figure 7 that there is a rather large vapor density gradient tangential to the surface of the inert particle. The vapor density gradient tangential to the surface of the ice particle is quite small. The presence of the inert particle significantly affects the profiles near the growing particle. As the particle separation decreases this effect becomes more pronounced, since the inert particle will occupy a greater solid angle, as measured from the center of the ice particle, from which vapor cannot be extracted.

B. DIFFUSIOPHORETIC FORCE

The diffusiophoretic force on each of two juxtaposed growing water droplets is zero. This is true since the velocity field which satisfies boundary conditions (5) and (6) as well as the boundary condition on the velocity far from the droplets, and which is a solution to the equation of motion can be written as:

$$\bar{v} = - \frac{D}{\rho_g} \nabla \rho_v .$$

As mentioned previously, this velocity field, to a very good approximation, can be interpreted as the gradient of a scalar function

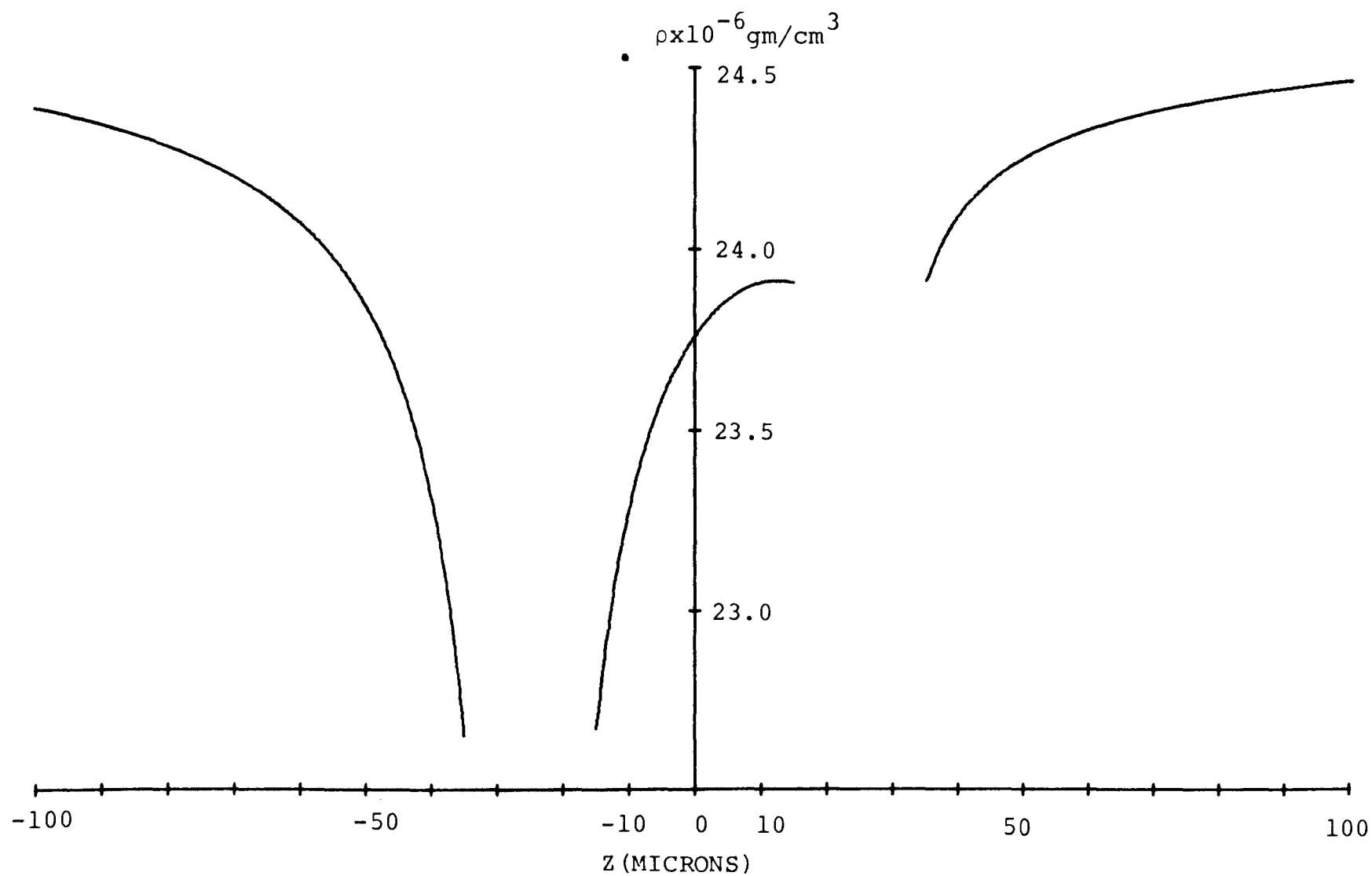


Figure 5. Vapor density profile for a juxtaposed water droplet (Right) and an ice particle. The supersaturation with respect to the water is 1.05 and with respect to the ice 1.135.

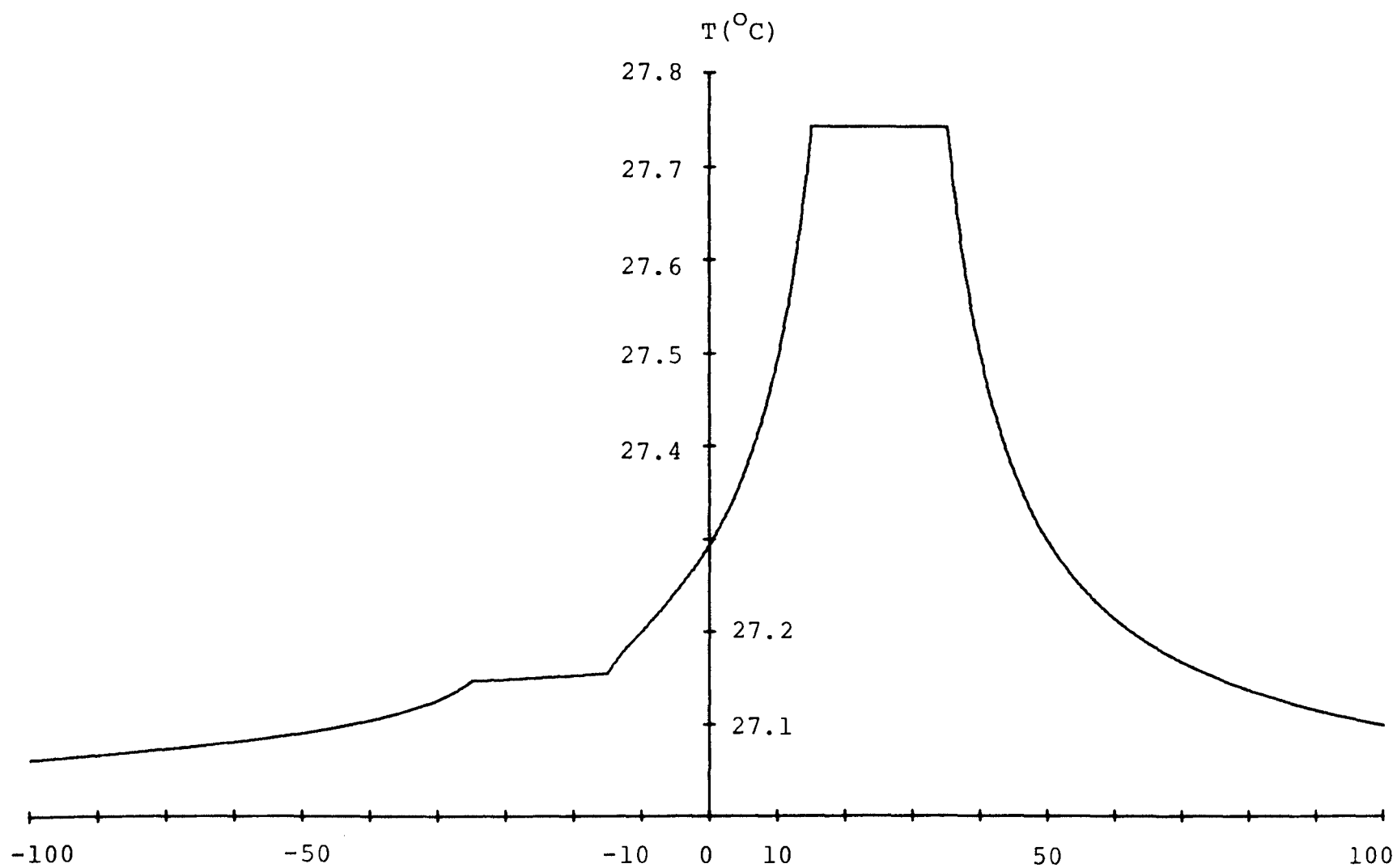


Figure 6. Temperature profile for a juxtaposed water droplet (right) and an ice particle. $S = 1.05$.

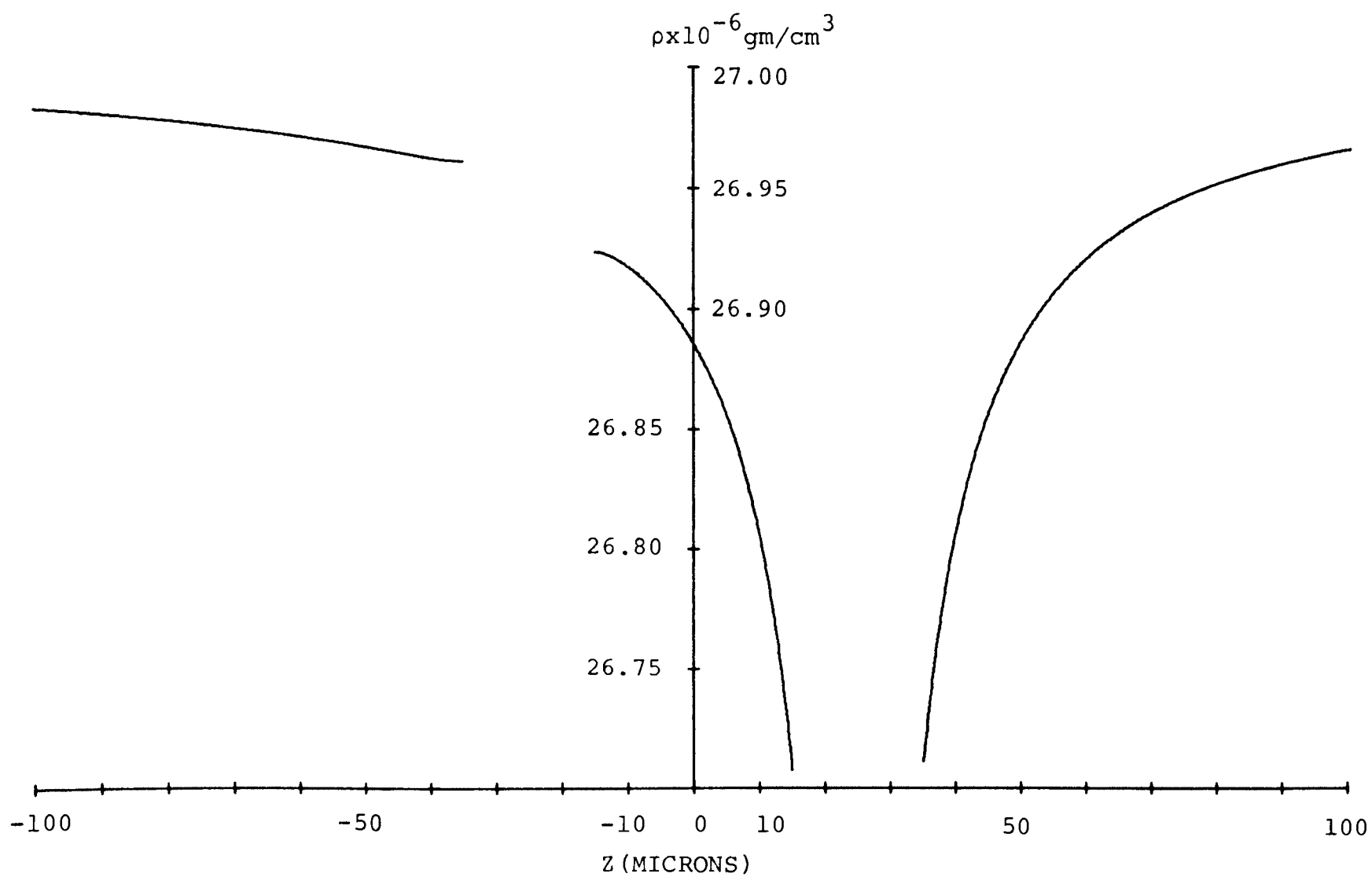


Figure 7. Vapor density profile for a juxtaposed water droplet (right) and an inert particle. $S = 1.05$.

and, hence, gives a zero force.

That this velocity field satisfies the velocity boundary condition tangential to the droplet surface is true because the internal droplet temperatures are constant. Conversely, if the internal droplet temperature of one of the droplets is not constant, the velocity field cannot be written as the gradient of a scalar unless the slip coefficient is taken as unity which is physically unacceptable.

The diffusiphoretic force on a spherical ice particle of radius 10 microns and an 8 micron water droplet for various separations is given in table I. The spheres are oriented such that the negative force on the ice particle means that it is attracted toward the water droplet. The force on the droplet is such as to repel it from the ice particle. Since the force on the droplet is larger, the overall relative force is repulsive. As the distance between surfaces increases from one to about ten microns the force on each particle decreases by more than an order of magnitude.

At lower supersaturations and small separations the situation may exist where the ice is growing and the water is evaporating. In this case the internal temperature gradient of the ice particle reverses in direction, hence the force reverses in direction. As a result, the force on each particle is such as to repel it from the other.

If the supersaturation is lowered to the point that both the ice and the water are evaporating, the force on the ice particle would be away from the water droplet, but the force on the water droplet would be toward the ice particle. Thus the overall relative

TABLE I

Diffusiophoretic force between an ice particle and a water droplet. F_1 is the force on the ice particle of radius 10 microns and F_2 is the force on the water droplet of radius 8 microns. D is the distance between centers. The supersaturation with respect to the droplet is 1.05 and with respect to the ice 1.137.

$D(\text{microns})$	$F_1(\text{dynes})$	$F_2(\text{dynes})$
0.1900E 02	-0.1184E 00	-0.1268E 00
0.1901E 02	-0.1164E 00	-0.1247E 00
0.1903E 02	-0.1137E 00	-0.1220E 00
0.1911E 02	-0.1038E 00	-0.1121E 00
0.1922E 02	-0.9307E-01	-0.1012E 00
0.1935E 02	-0.8208E-01	-0.9011E-01
0.1952E 02	-0.7145E-01	-0.7932E-01
0.1972E 02	-0.6152E-01	-0.6921E-01
0.1997E 02	-0.5251E-01	-0.6000E-01
0.2026E 02	-0.4449E-01	-0.5177E-01
0.2060E 02	-0.3746E-01	-0.4451E-01
0.2099E 02	-0.3138E-01	-0.3818E-01
0.2146E 02	-0.2616E-01	-0.3271E-01
0.2199E 02	-0.2173E-01	-0.2800E-01
0.2261E 02	-0.1799E-01	-0.2398E-01
0.2331E 02	-0.1484E-01	-0.2055E-01
0.2496E 02	-0.1012E-01	-0.1526E-01
0.2700E 02	-0.6874E-02	-0.1143E-01
0.2946E 02	-0.4648E-02	-0.8648E-02
0.3538E 02	-0.2210E-02	-0.5235E-02
0.4298E 02	-0.1057E-02	-0.3278E-02
0.5216E 02	-0.5178E-03	-0.2132E-02
0.6268E 02	-0.2618E-03	-0.1442E-02
0.7429E 02	-0.1366E-03	-0.1014E-02
0.8673E 02	-0.7314E-04	-0.7396E-03
0.9985E 02	-0.3955E-04	-0.5562E-03

force for situations in which both particles are evaporating would be attractive, since the magnitude of the force on the droplet would still be larger.

The forces on a spherical ice particle and a spherical inert particle for various separations are shown in table II. For this case, the forces are larger than the corresponding values in table I by about two orders of magnitude. The force on the inert particle is larger than that on the ice particle and is toward the ice particle. The force on the ice particle is away from the inert particle. The overall relative force is thus attractive. If the ice particle were evaporating the force on the inert particle would be such as to repel it from the ice making the overall relative force for this case repulsive.

Table III gives the force on a growing water droplet juxtaposed with an inert particle. Qualitatively, these results are identical to the situation shown in table II. Here the relative force is attractive and smaller than if the droplet were replaced by an ice particle, suggesting that ice is a more efficient scavenger than water.

That the diffusiophoretic force on a particle near an ice particle is larger than that force near a water droplet is true because the supersaturation of the bulk medium with respect to the ice surface is considerably larger than the supersaturation as seen from the water surface. Therefore, the flow of vapor toward the ice particle is larger than that toward the water droplet. The Stefan flow near the ice particle and the resulting force on an

TABLE II

Diffusiophoretic force between an ice particle and an inert particle. F_1 is the force on the 8 micron inert particle and F_2 the force on the 10 micron ice particle. D is the distance between centers. The supersaturation with respect to the ice particle is 1.137.

D(microns)	F_1 (dynes)	F_2 (dynes)
0.1900E 02	0.1036E 02	0.1147E 02
0.1902E 02	0.1019E 02	0.1130E 02
0.1906E 02	0.1000E 02	0.1111E 02
0.1917E 02	0.9430E 01	0.1053E 02
0.1931E 02	0.8787E 01	0.9882E 01
0.1949E 02	0.8097E 01	0.9182E 01
0.1970E 02	0.7378E 01	0.8453E 01
0.1995E 02	0.6651E 01	0.7713E 01
0.2024E 02	0.5933E 01	0.6981E 01
0.2059E 02	0.5238E 01	0.6270E 01
0.2100E 02	0.4581E 01	0.5593E 01
0.2147E 02	0.3969E 01	0.4960E 01
0.2201E 02	0.3411E 01	0.4376E 01
0.2264E 02	0.2909E 01	0.3846E 01
0.2335E 02	0.2464E 01	0.3369E 01
0.2416E 02	0.2075E 01	0.2946E 01
0.2507E 02	0.1738E 01	0.2572E 01
0.2609E 02	0.1451E 01	0.2244E 01
0.2838E 02	0.1012E 01	0.1723E 01
0.3112E 02	0.7029E 00	0.1328E 01
0.3433E 02	0.4877E 00	0.1029E 01
0.4169E 02	0.2465E 00	0.6460E 00
0.5067E 02	0.1290E 00	0.4181E 00
0.6102E 02	0.7106E-01	0.2806E 00
0.7249E 02	0.4136E-01	0.1956E 00
0.8483E 02	0.2540E-01	0.1414E 00
0.9787E 02	0.1637E-01	0.1055E 00

TABLE III

Diffusiophoretic force between a water droplet and an inert particle.
 F_1 is the force on the 10 micron water droplet and F_2 is the force
 on the 8 micron inert particle. D is the distance between centers.
 The supersaturation with respect to the droplet is 1.05.

$D(\text{microns})$	$F_1(\text{dynes})$	$F_2(\text{dynes})$
0.1900E 02	0.4236E 01	0.4665E 01
0.1902E 02	0.4185E 01	0.4613E 01
0.1904E 02	0.4121E 01	0.4549E 01
0.1914E 02	0.3906E 01	0.4331E 01
0.1926E 02	0.3658E 01	0.4081E 01
0.1941E 02	0.3387E 01	0.3807E 01
0.1959E 02	0.3100E 01	0.3517E 01
0.1981E 02	0.2808E 01	0.3220E 01
0.2008E 02	0.2516E 01	0.2923E 01
0.2039E 02	0.2231E 01	0.2633E 01
0.2076E 02	0.1960E 01	0.2355E 01
0.2119E 02	0.1705E 01	0.2093E 01
0.2168E 02	0.1471E 01	0.1850E 01
0.2225E 02	0.1259E 01	0.1628E 01
0.2291E 02	0.1070E 01	0.1427E 01
0.2366E 02	0.9038E 00	0.1249E 01
0.2450E 02	0.7592E 00	0.1090E 01
0.2644E 02	0.5342E 00	0.8362E 00
0.2880E 02	0.3723E 00	0.6422E 00
0.3161E 02	0.2583E 00	0.4953E 00
0.3820E 02	0.1296E 00	0.3077E 00
0.4647E 02	0.6665E-01	0.1963E 00
0.5623E 02	0.3584E-01	0.1296E 00
0.6722E 02	0.2035E-01	0.8891E-01
0.7919E 02	0.1220E-01	0.6326E-01
0.9193E 02	0.7705E-02	0.4657E-01
0.1052E 03	0.5087E-02	0.3532E-01

inert particle is greater than if the ice is replaced by a droplet.

C. FALL TRAJECTORIES

The significance of the diffusiophoretic force in terms of atmospheric processes for the particle sizes considered here can best be determined by comparing the particle trajectories with and without accounting for the diffusiophoretic force. To this end, table IV gives values of orientation and corresponding velocities for a 10 micron sphere falling toward an 8 micron sphere. The initial position was such that the larger sphere would come to within 1 micron separation between surfaces at the point of closest approach. Due to the interference of the flow fields around each sphere, which becomes more pronounced for small separations there is a component of horizontal velocity denoted as V_{1X} and V_{2X} , for each particle, giving the impression that the particles are "sliding along their line of centers." As noted experimentally by Eveson and Hall.⁴³ As the separation becomes smaller, the velocity of the smaller particle in the horizontal direction increases, in effect moving aside for the larger particle to pass. The vertical velocity of each sphere is larger than the Stokes value for isolated spheres due to the interference of the flow fields. For large sphere separations the fall velocity of each sphere approaches the Stokes value.

The diffusiophoretic force as discussed here is associated with diffusional growth or evaporation processes. A calculation of the trajectories to be accurate should account for changes in radius of

TABLE IV

Values of orientation and instantaneous velocity for two inert spheres.

D = distance between centers.

$\alpha = \tan^{-1} \frac{\Delta X}{\Delta Z}$ where ΔX and ΔZ are the horizontal and vertical separations respectively of the spheres.

V1X, V1Z = instantaneous velocities of sphere 1 (10 microns) in the X(horizontal) and Z(vertical) directions respectively.

V2X, V2Z = instantaneous velocities of sphere 2(10 microns) in the X(horizontal) and Z(vertical) directions respectively.

D(microns)	α (degrees)	V1X(^{cm} /sec)	V1Z(^{cm} /sec)	V2X(^{cm} /sec)	V2Z(^{cm} /sec)
0.1900E 02	0.900E 02	0	-0.1407E 01	0	-0.1258E 01
0.1901E 02	0.877E 02	-0.5896E-02	-0.1408E 01	-0.1064E-01	-0.1258E 01
0.1903E 02	0.854E 02	-0.1159E-01	-0.1409E 01	-0.2111E-01	-0.1259E 01
0.1911E 02	0.809E 02	-0.2175E-01	-0.1413E 01	-0.4079E-01	-0.1263E 01
0.1922E 02	0.764E 02	-0.3032E-01	-0.1418E 01	-0.5871E-01	-0.1268E 01
0.1936E 02	0.719E 02	-0.3735E-01	-0.1424E 01	-0.7461E-01	-0.1275E 01
0.1952E 02	0.674E 02	-0.4301E-01	-0.1430E 01	-0.8834E-01	-0.1284E 01
0.1973E 02	0.630E 02	-0.4751E-01	-0.1436E 01	-0.9982E-01	-0.1292E 01
0.1997E 02	0.586E 02	-0.5098E-01	-0.1442E 01	-0.1089E 00	-0.1302E 01
0.2026E 02	0.544E 02	-0.5351E-01	-0.1447E 01	-0.1156E 00	-0.1311E 01
0.2060E 02	0.503E 02	-0.5512E-01	-0.1452E 01	-0.1200E 00	-0.1319E 01

Con't next page

Table IV Con't.

D(microns)	α (degrees)	V1X($\frac{\text{cm}}{\text{sec}}$)	V1Z($\frac{\text{cm}}{\text{sec}}$)	V2X($\frac{\text{cm}}{\text{sec}}$)	V2Z($\frac{\text{cm}}{\text{sec}}$)
0.2100E 02	0.463E 02	-0.5583E-01	-0.1457E 01	-0.1219E 00	-0.1327E 01
0.2146E 02	0.426E 02	-0.5571E-01	-0.1460E 01	-0.1217E 01	-0.1332E 01
0.2200E 02	0.391E 02	-0.5482E-01	-0.1463E 01	-0.1196E 00	-0.1336E 01
0.2261E 02	0.358E 02	-0.5328E-01	-0.1465E 01	-0.1159E 00	-0.1337E 01
0.2332E 02	0.327E 02	-0.5120E-01	-0.1465E 01	-0.1109E 00	-0.1336E 01
0.2497E 02	0.272E 02	-0.4590E-01	-0.1463E 01	-0.9827E-01	-0.1326E 01
0.2701E 02	0.225E 02	-0.3983E-01	-0.1456E 01	-0.8410E-01	-0.1308E 01
0.2948E 02	0.187E 02	-0.3370E-01	-0.1445E 01	-0.7012E-01	-0.1281E 01
0.3539E 02	0.131E 02	-0.2283E-01	-0.1418E 01	-0.4639E-01	-0.1219E 01
0.4300E 02	0.948E 01	-0.1499E-01	-0.1386E 01	-0.2995E-01	-0.1152E 01
0.5219E 02	0.713E 01	-0.9863E-02	-0.1356E 01	-0.1951E-01	-0.1090E 01
0.6271E 02	0.558E 01	-0.6644E-02	-0.1330E 01	-0.1307E-01	-0.1038E 01
0.7432E 02	0.450E 01	-0.4624E-02	-0.1309E 01	-0.9075E-02	-0.9968E 00
0.8677E 02	0.374E 01	-0.3331E-02	-0.1292E 01	-0.6525E-02	-0.9639E 00
0.9988E 02	0.317E 01	-0.2477E-02	-0.1279E 01	-0.4848E-02	-0.9378E 00

growing or evaporating particles. This effect on the trajectories will be neglected, so that any differences in trajectories will correspond to diffusiophoretic forces.

Table V gives the relative orientation and velocities for a spherical ice particle of radius 10 microns falling toward a growing water droplet of radius 8 microns. Again, the initial positions are such that when the particle centers lie on the same horizontal plane, the surfaces are separated by 1 micron. It is apparent from comparison of the corresponding instantaneous velocities in tables IV and V that the effect of diffusiophoresis on the motion is not appreciable.

Values of relative orientation and instantaneous velocities for a 10 micron spherical ice particle and an 8 micron inert particle are shown in table VI. Comparing these values to the corresponding values in table IV, it appears that the diffusiophoretic force does noticeably influence the trajectories. The effect is most apparent for orientations in which the particles are positioned near the same horizontal plane, corresponding to $\alpha = 90^\circ$. For this orientation the horizontal force, derived from the interference of the flow fields around two settling spheres, is at its smaller values and due to the close proximity of the spheres the diffusiophoretic force is at its maximum value.

A plot of the relative position of the particles during the trajectory for the ice and inert particle interaction is shown in curve B of figure 8. For comparison, the corresponding plot for two inert spheres is given in curve A. Defining an impact parameter,

TABLE V

Values of orientation and instantaneous velocity for a 10 micron ice particle (sphere 1) falling toward an 8 micron water droplet. (See Table IV for explanation of symbols.) The supersaturation with respect to the water is 1.05 and with respect to the ice 1.137.

D(microns)	α (degrees)	V1X($\frac{\text{cm}}{\text{sec}}$)	V1Z($\frac{\text{cm}}{\text{sec}}$)	V2X($\frac{\text{cm}}{\text{sec}}$)	V2Z($\frac{\text{cm}}{\text{sec}}$)
0.1900E 02	0.900E 02	-0.6000E-03	-0.1407E 01	-0.6309E-03	-0.1258E 01
0.1901E 02	0.877E 02	-0.6485E-02	-0.1408E 01	-0.1126E-01	-0.1258E 01
0.1903E 02	0.854E 02	-0.1217E-01	-0.1409E 01	-0.2172E-01	-0.1259E 01
0.1911E 02	0.809E 02	-0.2228E-01	-0.1413E 01	-0.4134E-01	-0.1263E 01
0.1922E 02	0.764E 02	-0.3079E-01	-0.1419E 01	-0.5920E-01	-0.1268E 01
0.1935E 02	0.719E 02	-0.3775E-01	-0.1424E 01	-0.7504E-01	-0.1275E 01
0.1952E 02	0.674E 02	-0.4335E-01	-0.1430E 01	-0.8871E-01	-0.1284E 01
0.1972E 02	0.630E 02	-0.4779E-01	-0.1436E 01	-0.1001E 00	-0.1293E 01
0.1997E 02	0.586E 02	-0.5121E-01	-0.1442E 01	-0.1092E 00	-0.1302E 01
0.2026E 02	0.544E 02	-0.5369E-01	-0.1447E 01	-0.1158E 00	-0.1311E 01
0.2060E 02	0.503E 02	-0.5526E-01	-0.1452E 01	-0.1201E 00	-0.1319E 01
0.2099E 02	0.463E 02	-0.5595E-01	-0.1457E 01	-0.1221E 00	-0.1327E 01
0.2146E 02	0.426E 02	-0.5580E-01	-0.1460E 01	-0.1218E 00	-0.1332E 01
0.2199E 02	0.391E 02	-0.5489E-01	-0.1463E 01	-0.1197E 00	-0.1336E 01
0.2261E 02	0.358E 02	-0.5333E-01	-0.1465E 01	-0.1160E 00	-0.1337E 01
0.2331E 02	0.327E 02	-0.5125E-01	-0.1465E 01	-0.1109E 00	-0.1336E 01
0.2496E 02	0.272E 02	-0.4592E-01	-0.1463E 01	-0.9831E-01	-0.1326E 01
0.2700E 02	0.225E 02	-0.3985E-01	-0.1456E 01	-0.8413E-01	-0.1308E 01
0.2946E 02	0.187E 02	-0.3371E-01	-0.1445E 01	-0.7015E-01	-0.1282E 01
0.3538E 02	0.131E 02	-0.2284E-01	-0.1418E 01	-0.4640E-01	-0.1219E 01
0.4298E 02	0.947E 01	-0.1499E-01	-0.1386E 01	-0.2996E-01	-0.1152E 01
0.5216E 02	0.713E 01	-0.9865E-02	-0.1356E 01	-0.1952E-01	-0.1090E 01
0.6268E 02	0.557E 01	-0.6644E-02	-0.1330E 01	-0.1807E-01	-0.1038E 01
0.7429E 02	0.450E 01	-0.4624E-02	-0.1309E 01	-0.9075E-02	-0.9969E 00
0.8673E 02	0.374E 01	-0.3330E-02	-0.1292E 01	-0.6525E-02	-0.9640E 00
0.9985E 02	0.317E 01	-0.2476E-02	-0.1279E 01	-0.4847E-02	-0.9379E 00

TABLE VI

Values of orientation and relative velocity for a 10 micron ice particle (sphere 1) falling toward an 8 micron inert particle. (See Table IV for an explanation of the symbols.) The supersaturation with respect to the ice particle is 1.137.

D(microns)	α (degrees)	V1X(^{cm} /sec)	V1Z(^{cm} /sec)	V2X(^{cm} /sec)	V2Z(^{cm} /sec)
0.1900E 02	0.9000E 02	0.5335E-01	-0.1407E 01	0.5627E-01	-0.1258E 01
0.1902E 02	0.8775E 02	0.4655E-01	-0.1406E 01	0.4474E-01	-0.1256E 01
0.1906E 02	0.8549E 02	0.3981E-01	-0.1406E 01	0.3321E-01	-0.1254E 01
0.1917E 02	0.8096E 02	0.2647E-01	-0.1406E 01	0.1026E-01	-0.1253E 01
0.1931E 02	0.7640E 02	0.1414E-01	-0.1408E 01	-0.1159E-01	-0.1255E 01
0.1949E 02	0.7183E 02	0.2809E-02	-0.1411E 01	-0.3198E-01	-0.1258E 01
0.1970E 02	0.6728E 02	-0.7520E-02	-0.1415E 01	-0.5055E-01	-0.1264E 01
0.1995E 02	0.6278E 02	-0.1680E-01	-0.1419E 01	-0.6694E-01	-0.1270E 01
0.2024E 02	0.5838E 02	-0.2493E-01	-0.1424E 01	-0.8081E-01	-0.1278E 01
0.2059E 02	0.5409E 02	-0.3179E-01	-0.1430E 01	-0.9190E-01	-0.1286E 01
0.2100E 02	0.4997E 02	-0.3727E-01	-0.1435E 01	-0.1000E 00	-0.1294E 01
0.2147E 02	0.4602E 02	-0.4134E-01	-0.1440E 01	-0.1053E 00	-0.1301E 01
0.2201E 02	0.4227E 02	-0.4405E-01	-0.1444E 01	-0.1079E 00	-0.1307E 01
0.2264E 02	0.3874E 02	-0.4550E-01	-0.1447E 01	-0.1081E 00	-0.1310E 01
0.2335E 02	0.3544E 02	-0.4585E-01	-0.1450E 01	-0.1062E 00	-0.1312E 01
0.2416E 02	0.3238E 02	-0.4529E-01	-0.1451E 01	-0.1027E 00	-0.1310E 01
0.2507E 02	0.2955E 02	-0.4398E-01	-0.1450E 01	-0.9795E-01	-0.1306E 01
0.2609E 02	0.2695E 02	-0.4212E-01	-0.1449E 01	-0.9227E-01	-0.1299E 01
0.2838E 02	0.2241E 02	-0.3728E-01	-0.1443E 01	-0.7948E-01	-0.1280E 01
0.3112E 02	0.1869E 02	-0.3188E-01	-0.1432E 01	-0.6650E-01	-0.1254E 01
0.3433E 02	0.1569E 02	-0.2661E-01	-0.1419E 01	-0.5454E-01	-0.1223E 01
0.4169E 02	0.1129E 02	-0.1782E-01	-0.1389E 01	-0.3573E-01	-0.1158E 01
0.5067E 02	0.8459E 01	-0.1180E-01	-0.1359E 01	-0.2336E-01	-0.1096E 01
0.6102E 02	0.6579E 01	-0.7954E-02	-0.1333E 01	-0.1563E-01	-0.1044E 01
0.7249E 02	0.5291E 01	-0.5521E-02	-0.1311E 01	-0.1081E-01	-0.1001E 01
0.8483E 02	0.4376E 01	-0.3963E-02	-0.1294E 01	-0.7746E-02	-0.9678E 00
0.9787E 02	0.3703E 01	-0.2936E-02	-0.1280E 01	-0.5733E-02	-0.9409E 00

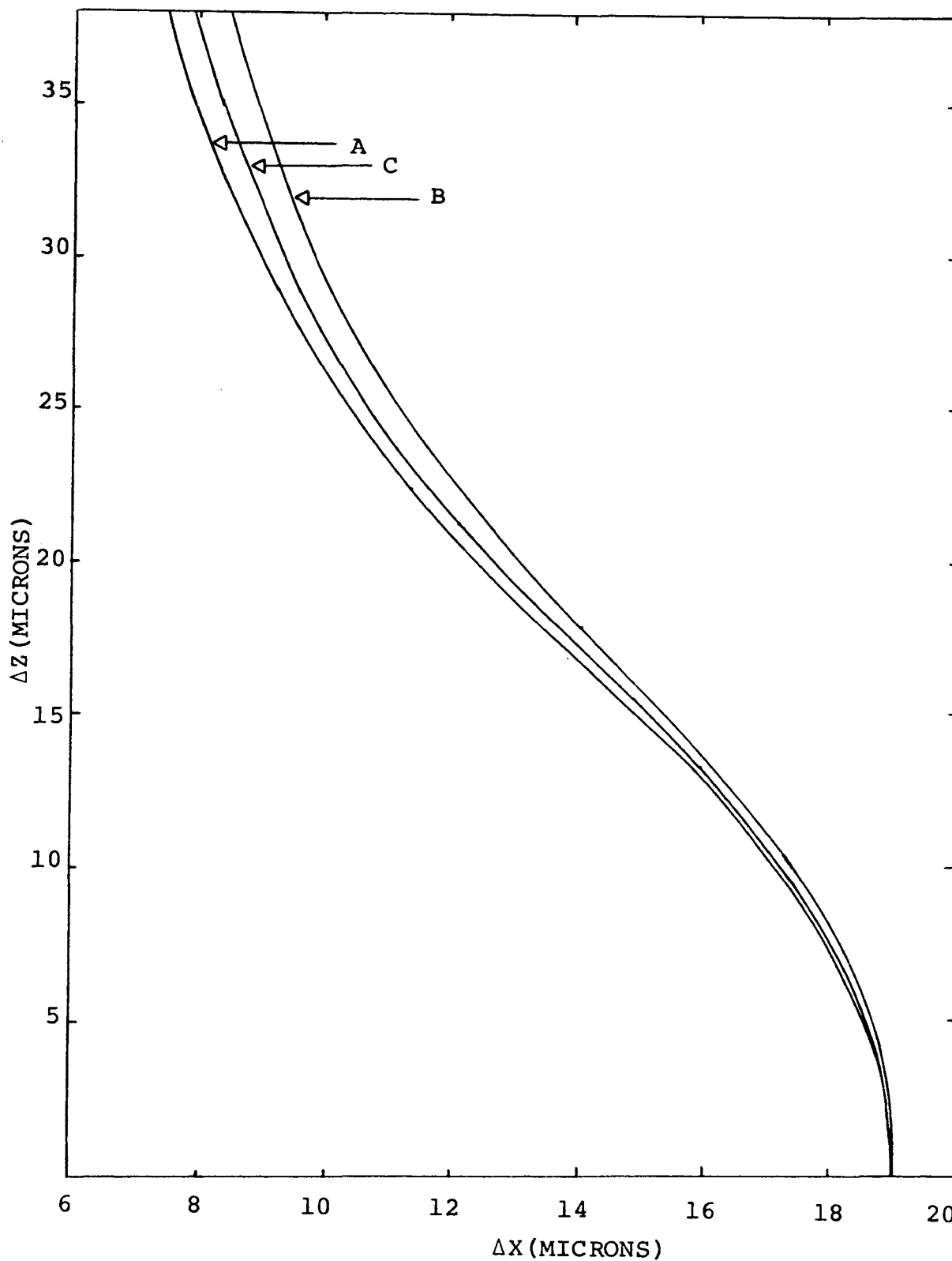


Figure 8. Fall trajectories for 8 and 10 micron spheres. Curve A represents two inert spheres, Curve B an ice and an inert particle and Curve C a water droplet and an inert particle. The supersaturation with respect to water is 1.05 and with respect to ice 1.137.

b, as the horizontal separation of the sphere centers at the beginning of a trajectory that results in a one micron miss, it appears that the diffusiophoretic force for a 10 micron ice particle and an 8 micron inert particle increases the impact parameter by about 1 micron.

Curve C of figure 8 gives the relative trajectory corresponding to a 10 micron growing water droplet falling toward an 8 micron inert particle and missing by 1 micron. The diffusiophoretic force does not change the trajectories as much for this situation as for the ice and inert particle trajectory. The phoretic force increases the impact parameter by about .5 microns.

Curve B of figure 9 shows the trajectory for a 10 micron growing ice particle and a 9.5 micron inert sphere. Shown in curve A is the corresponding trajectory for two inert spheres. Since the spheres are more nearly equal in size than the situation shown in figure 8, the relative fall velocity is smaller and the spheres spend more time in close proximity. As a result the phoretic force has a greater influence on the motion. For this situation the impact parameter is increased by 1.5 microns due to diffusiophoresis.

Recent calculations for collision efficiencies of small spheres falling under the assumptions of creeping, steady motion are given by Sartor and Davis³⁶ (1967). They defined the collision efficiency as $E = b^2 / (R_1 + R_2)^2$ where b is the impact parameter for a grazing collision and R_1 , R_2 the radii of the spheres. As mentioned earlier, the sphere surfaces are not permitted to touch in the present calculation. (The same criticisms can, of course, be made concerning

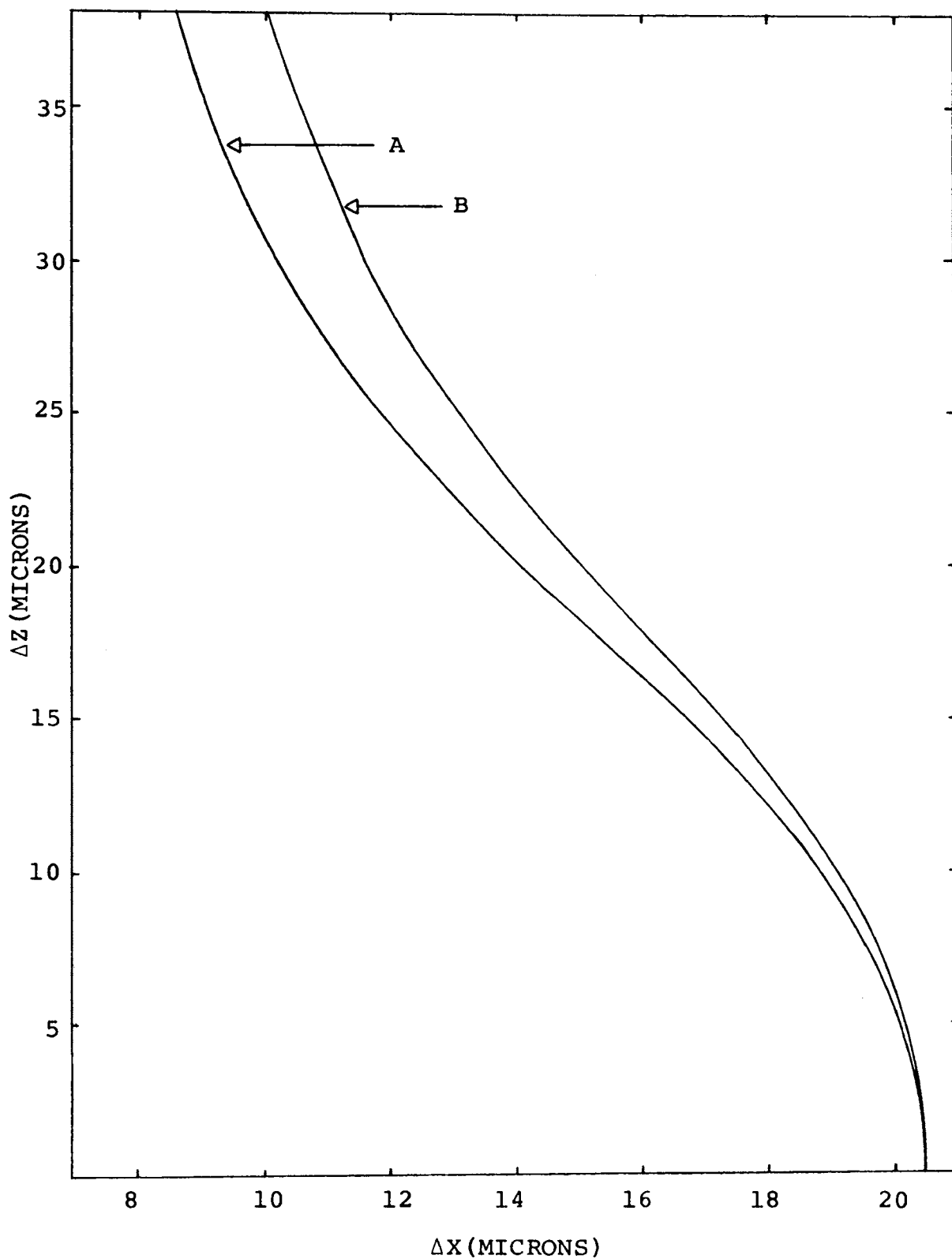


Figure 9. Fall trajectories for 9.5 and 10 micron spheres. Curve A corresponds to two inert spheres and curve B to an ice particle and an inert sphere. The supersaturation with respect to the ice is 1.135.

the Sartor and Davis calculations.) However, using the values of collision efficiencies given by Sartor and Davis and assuming that the changes in the impact parameters mentioned above due to diffusiophoresis apply to a grazing collision as well, one can get an idea of how diffusive forces change the impact parameters.

In the spirit of these assumptions, the impact parameter for a growing 10 micron ice sphere falling toward an 8 micron inert particle is increased by a factor of 1.4 due to diffusive forces. If the inert particle is increased in size to 9.5 microns the diffusive forces increase the impact parameter by a factor of 1.7.

Judging from curve C of figure 8 it appears that diffusive forces for the case of a 10 micron growing water droplet falling toward an 8 micron inert particle increases the impact parameter by a factor of 1.2.

V. APPLICATIONS AND SUGGESTIONS FOR FURTHER STUDY

Although the question as to what value to use for the diffusion slip coefficient is not known precisely, it does not appear to be a limiting factor in judging the significance of the phoretic force. The value for the slip coefficient as used in the tangential velocity boundary condition, equation (27), was $\sigma = .2$. For $\sigma = 1$ the force is zero. This corresponds to perfect slip at the particle surface. In the case of $\sigma = 0$ the force attains its maximum value.

A. ACCRETION AND IN-CLOUD SCAVENGING

The possibility of employing the above calculation for explaining problems involving cloud development does not look encouraging. Drop-let growth due to accretion is not affected at all by the phoretic force. Even the number of collisions of water droplets with ice particles do not appear to be significantly enhanced by diffusio-phoresis. These results are borne out by field observations of Rosinski⁵⁹ (1967) who concluded that "9 - 11 micron diameter cloud droplets will not be removed effectively by phoretic forces."

There does not appear to be sufficient experimental evidence to determine the effect of diffusiophoretic forces on the scavenging of inert particles of several microns radius by droplets or ice particles. The small amount of data presented by Rosinski⁵⁹ shows that the concentration of 9-11 micron diameter particles in air is about one-tenth and in ice about one-hundredth of that of 1.5 - 3 micron diameter particles. This data also indicates that the removal of 9-11 micron

diameter particles is independent of the type of ice phase. (One ice phase results from the accretion of supercooled droplets, the other from growth by sublimation.) Since the data points are scattered and include coagulation probabilities in addition to collision efficiencies, the above predictions concerning the enhancement of collision efficiencies of inert particles cannot be tested.

A very difficult problem related to the present discussion is that of the sticking together or coagulation of the particles after collision. If a diffusiophoretic force exists between particles, it may have an influence on the coagulation problem as well as the trajectories. Unfortunately, the diffusiophoretic forces calculated here are not reliable for very close sphere separations. However, these results do suggest that the phoretic force is large for small separations. For cases when the force is attractive, it may increase the chances of coagulation after the spheres collide.

The present study indicates that the diffusiophoretic force may enhance collisions between growing cloud elements and inert particles. The above results could be used to determine the extent of this increase. Such a calculation should necessarily take into account the number distributions of growing cloud elements and inert particles as well as the corresponding size distributions. The number distribution of inert particles is particularly complicated, since the possibility exists that any such particle may serve as a site of condensation of water vapor, thus becoming, in essence, a water droplet or ice particle.

The calculations presented here for the temperature and vapor

density profiles were done assuming that the velocity of each sphere relative to the surrounding medium was small enough that ventilation did not affect the density profiles. For particles of radius greater than about 20 microns, the fall velocity can affect the profiles.⁶⁰ Such a drop falling in a supersaturated atmosphere would experience an increase in the vapor density gradient on the front side. This increase in the gradient would increase the diffusiophoretic force on any particle that may be in the path of the falling drop. The result could be an enhancement of the collision efficiency.

Another very important item concerning the calculations presented here is the value of the bulk temperature of the medium. The equilibrium vapor pressure over ice at -12°C is much lower than that over water at the same temperature making the supersaturation over ice much larger than that over water. The difference between vapor pressures over ice and supercooled water is a function of temperature, being zero at 0°C and increasing to a maximum at -12°C .⁶¹ The difference in vapor pressures decreases almost linearly from -15°C to near -40°C . (At this temperature most of the supercooled droplets would be frozen.) There is, therefore, a limited temperature range in which the ice particles are more efficient at attracting inert particles than water droplets.

B. LIMITATIONS OF BISPHERICAL GEOMETRY

Another situation that can possibly be exploited by the calculations presented here is the diffusiophoretic force between ice crystals (not spherical) and either droplets or inert particles.

This suggestion is based on calculations by Podzimek²⁷ which show that the density profile around an ice crystal of complicated shape has nearly the same density profile as around a sphere. This is true since at large distances from the ice crystal the vapor density must necessarily reach the same bulk value as around the spherical ice particle. Also at close distances, any sharp point on the ice crystal would grow faster than a flat surface thus having a large effect on the profile in that region. Since the density profile around a sharp point has nearly spherical geometry, the density profile in the neighborhood of the ice crystal may be very similar to that around an ice sphere.

The effect of the diffusiophoretic force between an ice crystal of complicated geometry and an inert particle or a water droplet should be judged from the effect of the phoretic force on the fall trajectory. Unfortunately the fall trajectory for juxtaposed spheres does not appear to be of much use in this calculation, since the flow field around an ice crystal depends to a large extent on its shape.

Another application of the results presented here could be to give a guideline as to which situations it is necessary to use the relatively complicated bispherical geometry. If one is interested in the temperature and vapor density profiles in the vicinity of two widely spaced droplets, it may be possible to assume that one droplet is suspended in a medium on which is imposed constant, external temperature and vapor density gradients as was done by

Dukhin and Deryagin¹² Accuracy of such a calculation or the diffusio-phoretic force resulting from the calculation could be checked using the above results. Depending upon the amount of error one is willing to accept, there is a particle separation greater than which this approximation would be very useful.

C. SCAVENGING OF SUB-MICRON PARTICLES

The calculations presented here for the temperature and vapor density profiles for an ice particle juxtaposed with a water droplet and their relative trajectories could prove useful for the description of the scavenging of sub-micron particles in mixed clouds. Vittori and Prodi²⁸ (1967) suggested that the capture efficiency of small aerosol particles by ice crystals would be enhanced by the presence of supercooled droplets since the particles would be driven away from the evaporating droplet and toward the growing ice crystal by diffusiophoresis. Rosinski³⁰ (1967) concluded from field experiments that there is indeed an enhanced scavenging mechanism in mixed clouds for particles of 1.5 - 3 microns in diameter. Slinn and Hales^{23,24} have suggested that the particles are scavenged by the water droplets, since for these small particles thermophoresis dominates diffusio-phoresis and is in the opposite direction.

If the aerosol particles are considered small enough that their presence does not affect the vapor density and temperature profiles calculated here for a juxtaposed ice particle and water droplet system, then these profiles could be used to calculate the motion of the aerosol particles. The settling motion of the ice

particle and water droplet due to gravity must be taken into account for such a calculation. One would also be wise to consider the distributions of the particles involved to see if such a three-particle interaction is statistically probable before undertaking the calculation.

VI CONCLUSIONS

The diffusiophoretic force between two juxtaposed spherical water droplets is zero. Therefore, the trajectories of the two droplets when falling in a gravitational field are not influenced by the diffusive forces.

The diffusiophoretic force between a spherical ice particle and a spherical water droplet each of which is undergoing diffusional growth or evaporation is non-zero. However, if the sphere sizes are such that ventilation does not disturb the vapor density profiles and the hydrodynamic assumptions of steady, creeping motion are good, the diffusive forces acting between the spheres are not large enough to significantly change the trajectories.

If one of the spheres is a water droplet or spherical ice particle growing by diffusion and the other sphere represents an inert particle, the fall trajectories are influenced in favor of collision. The effect is greatest for the situation in which the spheres are of nearly equal size, since the relative velocities are then small and the spheres can spend more time in close proximity. The magnitude of the force depends on the supersaturation, being larger for higher supersaturations.

VII BIBLIOGRAPHY

1. Davis, M. H. "Collisions of Very Small Cloud Droplets," J. of Geophysical Research, 71, 3101 (1966).
2. Stefan, J., Wein Ber., 83, 943 (1881).
3. Hidy, G. M. and Brock, J. R., The Dynamics of Aerocollidal Systems, Pergamon Press Inc. (1970).
4. Kramers, H. A. and Kistemaker, J., "On the Slip of a Diffusing Gas Mixture Along a Wall," Physica, 10, 699 (1943).
5. Atkin, J., Transactions of the Royal Society Edinb., 32 (1883).
6. Watson, H., Transactions of the Faraday Society, 32, 1073 (1936).
7. Facy, L., Arch. Metero. Geophys. Bioklim., 8A, 269 (1955).
8. Facy, L., C. R. Acad. Sci., Paris, 246, 102, 3161 (1958).
9. Facy, L., Geof. Pura e Appl., 40, No. 2, (1958).
10. Deryagin, B. and Dukhin, S., Dokl. Akad. Nauk SSSR, 111, 613 (1956).
11. Deryagin, B. and Dukhin, S., Dokl. Akad. Nauk SSSR, 106, 851 (1956).
12. Dukhin, S. and Deryagin, B., "A Theory of Interaction between Evaporating or Growing Droplets at Large Distances," Dokl. Akad. Nauk SSSR, 112, 407 [Soviet Physics Doklady, 2, 41 (1957) (English Translation)].
13. Waldmann, L., "On the Motion of Spherical Particles in Non-Homogeneous Gases" in Rarefied Gas Dynamics, ed. L. Talbot, Academic Press, New York, 323 (1961).
14. Waldmann, L. and Schmitt, K. H., "Thermophoresis and Diffusiophoresis of Aerosols" in Aerosol Science, ed. C. N. Davies, Academic Press, New York (1967).
15. Schmitt, K. H., and Waldmann, L., Z Naturforschg, 15d, 843 (1960).
16. Schmitt, K. H., "Untersudnungen an Schwebstoffteilchen in Diffundierendem Wasserdampf," Z. Naturforschg., 16a, 144 (1961).
17. Brock, J. R., "Forces on Aerosols in Gas Mixtures," J. of Colloid Science, 18, 489-501 (1963).

18. Deryagin, B. V., Yalamov, Y. I. and Storozhilova, A. I., "Diffusiophoresis of Large Aerosol Particles," J. of Colloid and Interface Science, 22, 117-125 (1966).
19. Deryagin, B. V. and Yalamov, "Theory of Motion of Large Volatile Particles in Diffusing Gas Mixtures, Soviet Physics - Doklady, 12, 587 (1967).
20. Deryagin, B. V. and Yalamov, "The Theory of the Motion of Solution Drops in a Diffusing Gas Mixture," Soviet Physics - Doklady, 12 (1968).
21. Prokhorov, P. S. and Lenov, L. F., "Investigation of Long - Range Diffusion Forces Between Water Droplets and Non-Volatile Particles," Discussions of the Faraday Society, 30, 124 (1960).
22. Waldmann, L., Z. Naturforschg., 30, 124-129 (1961).
23. Slinn, W. G. N. and Hales, J. M., "Phoretic Forces in Scavenging," Proceedings of a Symposium on Precipitation Scavenging, Richland, Washington, June 2 - 9, 1970, National Technical Information Service, U.S. Dept. of Commerce, CONF. - 700601.
24. Slinn, W. G. N. and Hales, J. M., "A Reevaluation of the Role of Thermophoresis as a Mechanism of In-and Below-Cloud Scavenging," J. Atmos. Sci. 8, 1465 (1971).
25. Goldsmith, P., Delafield, H. J and Cox, L. C., "The Role of Diffusiophoresis in the Scavenging of Radioactive Particles From the Atmosphere," Q. J. R. Met. Soc., 89, 43 (1963).
26. Goldsmith, P. and May, F. G., "Diffusio-and Thermophoresis in Water Vapor Systems," in Aerosol Science, ed. C. N. Davies, Academic Press, New York, (1966).
27. Podzimek, J., "L'Influence Du Courant De Stefan Sur La Capture Des Particules D'AgI Sur Les Elements De Nuage," Journal De Recherches Atmospheriques, 19 (1965).
28. Vittori, O. A., and Prodi, V., "Scavenging of Atmospheric Particles by Ice Crystals," J. Atmos. Sci., 24, 533 (1967).
29. Vittori, O. A. and Prodi, V., "Particle-Free Space in Stefan Flow," Proceedings of a Symposium on Precipitation Scavenging, Richland, Washington, June 2-9, 1970, National Technical Information Service, U.S. Dept. of Commerce, CONF-700601.
30. Rosinski, J., "Insoluble Particles in Hail and Rain," J. Applied Meteorolgy, 6, 1066 (1967).
31. Hocking, L. M., Ph.D Dissertation, University of London, (1958) Unpublished.

32. Stimson, M., and Jeffery, G. B., "The Motion of Two Spheres in a Viscous Fluid," Proc. Roy. Soc. London, A, 111, 110 (1926).
33. Maude, A. D., "End Effects in a Falling-sphere Viscometer," Brit. J. App. Physics, 12, 293 (1961).
34. Brenner, H., "The Slow Motion of a Sphere Through a Viscous Fluid Towards a Plane Surface," Chem. Eng. Sci., 16, 242 (1961).
35. Wakiya, S., Niigata Univ. (Nagooka Japan) Coll. Eng. Res. Report, No. 6, March 30 (1957).
36. Davis, M. H. and Sartor, J. D., "Theoretical Collision Efficiencies for Small Cloud Droplets in Stokes Flow," Nature, 215, 1371 (1967).
37. O'Neill, M. E., "Slow Motion of Viscous Liquid Caused by a Slowly Moving Solid Sphere," Mathematika, 11, 67 (1964).
38. Dean, W. R. and O'Neill, M. E., "A Slow Motion of Viscous Liquid Caused by the Rotation of a Solid Sphere," Mathematika, 10, 13 (1963).
39. Wakiya, S., "Slow Motions of a Viscous Fluid Around Two Spheres," J. Phys. Soc. Japan, 22, 1101 (1967).
40. Davis, M. H., "The Slow Translation and Rotation of Two Unequal Spheres in a Viscous Fluid," Chemical Engineering Science, 24, 1769 (1969).
41. Happel, J. and Brenner, H., Low Reynolds Number Hydrodynamics, Englewood Cliffs, New Jersey, Prentice-Hall, 516 (1965).
42. Goldman, A. J., Cox, R. G., and Brenner, H., "The Slow Motion of Two Identical Arbitrarily Oriented Spheres Through a Viscous Fluid," Chem. Eng. Sci., 21, 1151 (1966).
43. Eveson, G. F., Hall, E. W., and Ward, S. G., "Interaction Between Two Equal-Sized Equal-Settling Spheres Moving Through a Viscous Liquid," British J. of Applied Physics, 10, 43 (1959).
44. Mathews, H. W., and Smith, F. B., Br. J. of Appl. Phys., 11, 87 (1960).
45. Slack, G. W., and Mathews, H. W., "The Mutual Interaction Between Sedimenting Particles, Part I. The Sedimentations of Compact Clusters of Spheres in a Viscous Medium," Chemical Defense Experimental Establishment, Porton, Nr. Salisbury (1961).
46. Jayawiera, K. O. L. F., Mason, B. J., and Slack, G. W., J Fluid Mech., 20, 121 (1964).

47. Klett, J. D., "The Interaction and Motion of Rigid Spheres Falling in a Viscous Fluid at Low Reynolds Numbers," Ph.D. Dissertation, U.C.L.A. (1968) Unpublished.
48. Landeau, L. D. and Lifshitz, E. M., Fluid Dynamics, Pergamon Press (1959).
49. Bird, R. B., Stewart, W. E., and Lightfoot, E. N., Transport Phenomena, New York, Wiley (1960).
50. Brock, J. R., "On the Theory of Thermal Forces Acting on Aerosol Particles," Journal of Colloid Science, 17, 768-780 (1962).
51. Brock, J. R., "Investigation of a First-Order Slip-Flow Continuum Analysis: The Thermal Force," Journal of Physical Chemistry, 66, 1763 (1962).
52. Berker, R. in Encyclopedia of Physics: Fluid Dynamics II Vol VIII/2, p. 247. Springer-Verlog, Berlin (1963).
53. Brenner, H., "The Stokes Resistance of an Arbitrary Particle - II An Extension," Chem. Eng. Sci., 19, 599 (1964).
54. Carstens, J. C. and Kassner, J. L. Jr., "Some Aspects of Droplet Growth Theory Applicable to Nuclei Measurements" J. Rech. Atmos., 3, 33 (1968).
55. Morse, P. M. and Feshbach, H., Methods of Theoretical Physics, McGraw-Hill, New York, 1298 (1953).
56. Carstens, J. C., Williams, A., and Zung, J., "Theory of Droplet Growth in Clouds: II Diffusional Interaction Between Two Growing Droplets," J. Atmos. Sci., 27, 798 (1970).
57. Williams, A. L., and Carstens, J. C., "A Note Concerning the Interaction of Two Growing Water Droplets." J. Atmos. Sci., 28, 1298 (1971).
58. Rainville, E. D., Special Functions, The Macmillan Company, New York (1960).
59. Rosinski, J., "A Possible Role of Ice Forming Nuclei in Rain Formation" J of Applied Meteor., 6, 1062 (1967).
60. Abraham, Farid F., "A Physical Interpretation of the Structure of the Ventilation coefficients for Freely Falling Waterdrops" Journal of Atmos. Sci., 25, 76 (1968).
61. Fleagle, R. G. and J. A. Businger, An Introduction to Atmospheric Physics, Academic Press, New York and London (1963).

VITA

The author was born August 19, 1947, in Effingham, Illinois. He graduated from North Clay Community High School in 1965 and entered Eastern Illinois University, where he majored in physics and received a BS. in Ed. in August 1968. Since entering graduate school at the University of Missouri at Rolla, in September 1968, he has received the M.S. in Physics in May 1970 and has been a doctoral candidate in the Graduate Center for Cloud Physics Research, University of Missouri at Rolla.

The author is married to the former Miss Janice Sue Logsdon.

APPENDIX

A. Computer Program

The first 198 lines of the program are mainly devoted to reading in the necessary data and setting the position of the two spheres in bispherical coordinates. Solutions to the boundary condition equations specifying the unknown constants for determination of the temperature and vapor density profiles are given in lines 198-509. By altering the index IR this part of the program can be made to describe either a juxtaposition of two volatile particles or the situation in which one particle is volatile the other inert.

Lines 509 - 597 are devoted to calculating the diffusiophoretic force and a check on this calculation is done in lines 598 - 611.

The sum of the gravitational force on each particle parallel to the line joining centers and the diffusiophoretic force, as well as the component of the gravitational force perpendicular to the line of centers is found in lines 614 - 619.

With the forces known, the instantaneous velocities of each particle parallel to the line of centers (lines 620 - 700) and the corresponding velocities for motion perpendicular to the line of centers (lines 701 - 1046) are found.

The particles are allowed to fall an increment of time, DT , at these velocities resulting in a new orientation for which the process of calculating the diffusiophoretic force and the resulting instantaneous velocities is repeated. The time increment, DT , is adjusted (lines 1047 - 1064) at various points in the trajectory

according to specified tolerances to keep both the error in the trajectory and the computer time at reasonably small values. (The approximate computer time for one trajectory is about one hour.)

SUBROUTINE ALW - The program is treated as a subroutine and stored on a disc to avoid wasting compile time.

A1,A2 - Radii of spheres 1 and 2 respectively.

X10,X20 - Horizontal positions of spheres 1 and 2.

Z10,Z20 - Vertical positions of spheres 1 and 2.

BC1,BC2,CC1,CC2 - Constants defined in the thermal equilibrium boundary condition.

AD - Diffusion coefficient.

CNLAT1,CNLAT2 - Latent heats of spheres 1 and 2.

AKO - Thermal conductivity of saturated air.

AKI1,AKI2 - Internal thermal conductivities of spheres 1 and 2 respectively

TR,RHR - Bulk values of temperature and vapor density respectively.

SS1,SS2 - Supersaturation with respect to spheres 1 and 2 respectively.

GRAV - Acceleration of gravity.

V1X,V2X - Horizontal velocities of spheres 1 and 2 respectively.

V1Z,V2Z - Vertical velocities of spheres 1 and 2.

DA - Distance between sphere centers.

ALF - Defined as α in text (Table 4).

U10 = μ_1

U20 = μ_2

A(N,M),B(N) - Coefficient matrix and resultant vector to solve temperature and density profile boundary condition equations.

AA(N,M),BB(N) - Coefficient matrix and resultant vector to solve boundary condition equations for diffusiophoretic force.

DGELG - Tabulated subprogram for solving simultaneous equations.

F1,F2 - Diffusiophoretic force on particles 1 and 2 respectively.

FPL1,FPL2 - Force parallel to the line joining centers on particles
1 and 2 respectively.

FPD1,FPD2 - Force perpendicular to the line of centers on particles
1 and 2.

```

0001      DOUBLE PRECISION AP(122,122),BP(122)
0002      93 FORMAT(4I6)
C      IR=1 DENOTES INERT PARTICLE  IR=2--- VOLATILE PARTICLE
C      IN=1 ACTIVATES PROFILE CALC. IN=2 DOES NOT
C      IP DENOTES DIMENSION OF PROFILE MATRIX
C      JF DENOTES DIMENSION OF PERPENDICULAR FORCE MATRIX
0003      READ(1,93)IP,JF,IR,IN
0004      CALL ALW(AP,BP,JF,IP,IR,IN)
0005      STOP
0006      END
0001      SUBROUTINE ALW(AP,BP,JF,IP,IR,IN)
C      ATMOSPHERIC CONDITIONS
0002      IMPLICIT REAL*8(A-H,O-Z)
0003      REAL*4 ETS
0004      DIMENSION AP(JF,JF),BP(JF),AN1(35),AN2(35),BN1(35),BN2(35)
0005      DIMENSION AB1(35),AB2(35),CD1(35),CD2(35),EF1(36),EF2(36),GH1(34)
0006      1, GH2(34)
0007      DIMENSION ARAP(35),ARBP(35),ARAM(35),ARBM(35)
0008      DIMENSION X(300)
0009      DIMENSION PHE1(20),PHE2(20),PHI1(20),PHI2(20),GAP1(20),GAP2(20),
0010      1H(20),E(20),C(20),D(20),AA(4,4),BB(4),A(96,96),B(96)
0011      DIMENSION SJ1(10),SJ2(10),SJ3(10),SJ4(10),DM1(10),DM2(10),DM3(10),
0012      1DM4(10),RE1(10),RE2(10),RE3(10),RE4(10)
0013      DIMENSION ASP(10),BSP(10),CSP(10),DSP(10)
0014      DIMENSION AS(10),BS(10),CS(10),DS(10)
0015      DIMENSION X1W(5),W1X(5)
0016      DIMENSION VE1X(120),VE2X(120),VE1Z(120),VE2Z(120),FOR1(120),
0017      1FOR2(120),DIST(120),ANGL(120)
0018      DIMENSION CHGX(120),CHGZ(120)
0019      DOUBLE PRECISION DSQRT,DLOG,DTANH,DEXP,DCOSH,DSINH,DABS
0020      DOUBLE PRECISION DATAN,DCOS,DSIN,DTAN
0021      1 FORMAT (4D18.8)
0022      100 FORMAT(3D18.8)
0023      2 FORMAT(//)
0024      3 FORMAT (6D16.6)
0025      5 FORMAT (E18.8,16)
0026      11 FORMAT(' THE RESULTANT VECTOR -- B(N) ')
0027      12 FORMAT('      Z              U              T
0028      1      RHO')
0029      13 FORMAT(' THE COFICIENTS OF TEMP + DENSITY ')
0030      14 FORMAT('      R1              R2              S1              S2
0031      1      S1E              S2E              ')
0032      15 FORMAT(' THE COEFFICIENT MATRIX IS -- A(N,J) ')
0033      16 FORMAT(T8,'A1',T24,'A2')
0034      17 FORMAT('      BC1              BC2              CC1
0035      1      CC2 ')

```

```

0029      18 FORMAT(' --- AD CNLAT1 CNLAT2
0030      19 FORMAT(' AKO ')) AKI1 AKI2 TR
0031      21 FORMAT(' GAME1 GAME2 GAM11
0032      22 FORMAT(' BP1 BP2 ')) BP3 BP4 CP1
0033      23 FORMAT(' CP2 TD1 TD2 '))
0034      24 FORMAT(T6,'DA',T22,'U10',T38,'U20',T54,'AM',T70,'DA1',T86,'DA2')
0035      25 FORMAT(' Z1R Z1L Z2R
0036      31 FORMAT(' Z2L ')) ETS T1 IER ')) Y2
0037      34 FORMAT(' Y1 T2 ')) VIS MG MV
0038      37 FORMAT(' SLP '))
0039      38 FORMAT(' RHO1 RHO2 '))
0040      39 FORMAT(' PHE1 PHE2 ')) C(N)
0041      40 FORMAT(' H(N) '))
0042      41 FORMAT(' D(N) '))
0043      42 FORMAT(' DA V1 V2 '))
0044      44 FORMAT(' AS BS CS
0045      45 FORMAT(' DS '))
0046      46 FORMAT(' AA CC B1 B2
0047      47 FORMAT(' D1 D2 '))
0048      52 FORMAT(' PHI1 PHI2 '))
0049      53 FORMAT(' GAP2 '))
0050      54 FORMAT(' ASP BSP CSP
0051      55 FORMAT(' DSP '))
0052      56 FORMAT(' Y1P Y2P Y1PP Y2PP
0053      57 FORMAT(' VP VPP '))
0054      58 FORMAT(' F1',T24,'F2')
0055      59 FORMAT(' SI1',T24,'SI2',T42,'DSI1',T60,'DSI2')
0056      60 FORMAT(' V1',T24,'V2')
0057      61 FORMAT(' THE PERPENDICULAR VELOCITY COEFFICIENTS ARE '))
0058      62 FORMAT(' COLLISION '))
0059      63 FORMAT(' HIT '))
0060      64 FORMAT(' MISS '))
0061      65 FORMAT(' DA',T24,'ALF',T40,'EFF',T56,'A1A2')
0062      66 FORMAT(' DA',T24,'ALF')
0063      67 FORMAT(' X10',T24,'X20')
0064      68 FORMAT(' VALUES FROM LOOP 701 '))
0065      69 FORMAT(' VALUES FROM LOOP 700 '))
0066      70 FORMAT(' DA',T24,'ALF',T40,'EFF')
0067      71 FORMAT(' DA',T24,'ALF',T40,'TT')
0068      72 FORMAT(' S AND J BOUNDARY CONDITION CHECK '))

```

```

0063      67 FORMAT(' MAUDE BOUNDARY CONDITION CHECK ')
0064      68 FORMAT(T8,'SJ1',T24,'SJ2',T40,'SJ3',T56,'SJ4')
0065      69 FORMAT(T8,'RE1',T24,'RE2',T40,'RE3',T56,'RE4')
0066      70 FORMAT(T8,'DM1',T24,'DM2',T40,'DM3',T56,'DM4')
0067      71 FORMAT(T8,'FPD1',T24,'FPD2',T40,'FPD1P',T56,'FPD2P')
0068      72 FORMAT(T8,'CHK1',T24,'CHK2')
0069      73 FORMAT(T8,'A1',T24,'A2')
0070      76 FORMAT(T8,'AN1',T24,'BN1',T40,'AN2',T56,'BN2')
0071      77 FORMAT(5D17.7)
0072      78 FORMAT(16)
0073      79 FORMAT(T8,'AB1',T24,'AB2',T40,'CD1',T50,'CD2')
0074      80 FORMAT(T8,'EF1',T24,'EF2')
0075      81 FORMAT(T8,'GH1',T24,'GH2')
0076      82 FORMAT(T8,'CND11',T24,'CND12')
0077      83 FORMAT(T8,'CND21',T24,'CND22')
0078      84 FORMAT(T8,'CND31',T24,'CND32')
0079      85 FORMAT(T8,'CND41',T24,'CND42')
0080      86 FORMAT(T8,'CND51',T24,'CND52')
0081      87 FORMAT(T8,'ARB1',T24,'ARB2',T39,'ARB3',T56,'BRB1',T74,'BRB2',T89
0082      1,'BRB3')
0083      88 FORMAT(' THE MATRIX ELEMENTS AFTER MULTIPLICATION ARE --- ')
0084      89 FORMAT(' SUM ')
0085      90 FORMAT(5D26.16)
0086      91 FORMAT(' CROSS OVER ')
0087      93 FORMAT(4I6)
0088      94 FORMAT(4D26.16)
0089      95 FORMAT(T8,'X1',T24,'X2',T40,'Z1',T56,'Z2')
0090      96 FORMAT(T8,'V1X',T24,'V2X',T40,'V1Z',T56,'V2Z')
0091      97 FORMAT(' DT=DT/2 ')
0092      98 FORMAT(' DT=DT*2 ')
0093      99 FORMAT(T7,'DA',T23,'ALF',T39,'V1X',T55,'V1Z',T71,'V2X',T87,'V2Z')
0094      75 FORMAT(T8,'DA',T25,'F1',T42,'F2')
0095      32 FORMAT(T8,'SS1',T24,'SS2',T42,'RHR',T56,'SLP')

0095      WRITE(3,93)IP,JF,IR,IN
0096      READ(1,1)A1,A2
0097      READ(1,1)X10,X20,Z10,Z20
0098      READ(1,1)BC2,BC1,CC2,CC1
0099      READ(1,1)AD,CNLAT2,CNLAT1,AKO
0100      READ(1,1)SLP,PI,RHG,VIS
0101      READ(1,1)AKI2,AKI1,TR,SS2
0102      READ(1,1)TOLA,TOLR,DALF,GRAV
0103      READ(1,1)ERR1,ERR2,DT
0104      JE=1
0105      RHR=SS2*(BC2*TR+CC2)
0106      SS1=RHR/(BC1*TR+CC1)
0107      RHOI=.917

```

```

0108 WRITE(3,16)
0109 WRITE(3,1)A1,A2
0110 WRITE(3,2)
0111 WRITE(3,17)
0112 WRITE(3,1)BC1,BC2,CC1,CC2
0113 WRITE(3,2)
0114 WRITE(3,18)
0115 WRITE(3,1)AD,CNLAT1,CNLAT2,AKO
0116 WRITE(3,2)
0117 WRITE(3,19)
0118 WRITE(3,1)AKI1,AKI2,TR,RHR
0119 WRITE(3,2)
0120 SQ2=DSQRT(2.DO)
0121 GAMI1=AKI1/(AD*CNLAT1)
0122 GAMI2=AKI2/(AD*CNLAT2)
0123 GAME1=AKO/(AD*CNLAT1)
0124 GAME2=AKO/(AD*CNLAT2)
0125 BP1=BC1+GAME1
0126 BP2=BC2+GAME2
0127 BP3=BC2+GAME1
0128 BP4=BC1+GAME2
0129 TS=TR*SQ2
0130 CP1=SQ2*(CC1-RHR)
0131 CP2=SQ2*(CC2-RHR)
0132 TO1=(RHR-CC1+GAME1*TR)/BP1
0133 TO2=(RHR-CC2+GAME2*TR)/BP2
0134 WRITE(3,21)
0135 WRITE(3,3)GAME1,GAME2,GAMI1,GAMI2,BP1,BP2
0136 WRITE(3,2)
0137 WRITE(3,22)
0138 WRITE(3,3)BP3,BP4,CP1,CP2,TO1,TO2
0139 WRITE(3,2)
0140 WRITE(3,32)
0141 WRITE(3,1)SS1,SS2,RHR
0142 WRITE(3,2)
0143 801 X1=X10
0144 V1X=0
0145 V2X=0
0146 V1Z=0
0147 V2Z=0
0148 TT=0.
0149 INDP=1
0150 IND=0
0151 X2=X20
0152 Z1=Z10

```

```

0153      Z2=Z20
0154      606 CONTINUE
0155      DZ=Z1-Z2
0156      DX=X1-X2
0157      ARUG=DX*DX+DZ*DZ
0158      DA=DSQRT(ARUG)
0159      IF (DZ) 831,832,831
0160      832 ALF=PI/2.
0161      GO TO 835
0162      831 ALF=DATAN(DX/DZ)
0163      835 EFF=DX*DX/((A1+A2)**2)
0164      WRITE(3,62)
0165      WRITE(3,64)
0166      WRITE(3,1)DA,ALF,EFF
0167      TRACE=1.
0168      IF (EFF-.0005)825,825,800
0169      800 SAM=2*DA
0170      DIST(JE)=DA*10000.
0171      ANGL(JE)=ALF*57.29578
0172      DA1=(DA*DA+A1*A1-A2*A2)/SAM
0173      DA2=(DA*DA+A2*A2-A1*A1)/SAM
0174      AM=DSQRT((DA2+A2)*(DA2-A2))
0175      U10=DLOG((DA1+AM)/(DA1-AM))/2.DO
0176      U20=DLOG((DA2+AM)/(DA2-AM))/2.DO
0177      WRITE(3,73)
0178      WRITE(3,1)A1,A2
0179      WRITE(3,23)
0180      WRITE(3,3)DA,U10,U20,AM,DA1,DA2
0181      WRITE(3,2)
0182      F1=0.
0183      F2=0.
0184      IND=IND+1
0185      IF (IND-4)600,601,601
0186      601 W1X(1)=W1X(2)
0187      W1X(2)=W1X(3)
0188      X1W(1)=X1W(2)
0189      X1W(2)=X1W(3)
0190      IND=3
0191      600 CONTINUE
0192      X1W(IND)=X1
0193      V1XN=V1X
0194      V2XN=V2X
0195      V1ZN=V1Z
0196      V2ZN=V2Z
0197      IF (IN-1)699,699,799
0198      699 CONTINUE
C      SOLUTION TO BOUNDARY CONDITION EQNS.

```

```

0199      I=IP
0200      I6=I/6
0201      I23=2*I/3
0202      I3=I/3
0203      I2=I/2
0204      I3P=I3+1
0205      I2P=I2+1
0206      I23P=I23+1
0207      I56=5*I/6
0208      I56M=I56-1
0209      IM2=I-2
0210      I23M=I23-1
0211      I2M=I2-1
0212      I6M=I6-1
0213      I6MM=I6M-1
0214      I6P=I6+1
0215      I3M=I3-1
0216      I56P=I56+1
0217      DO 700 N=1,I
0218      DO 700 J=1,I
0219      700 A(N,J)=0.0
0220      DO 701 N=1,I23
0221      J=N
0222      701 A(N,J)=1.0
0223      DO 702 N=1,I6
0224      J=N+I6
0225      K=N-1
0226      CK=K
0227      U1K=(2.*CK+1.)*U10
0228      702 A(N,J)=DEXP(U1K)
0229      DO 703 N=1,I3
0230      J=N+I23
0231      703 A(N,J)=-1.0
0232      DO 704 J=1,I6
0233      N=J+I6
0234      K=J-1
0235      CK=K
0236      U2K=(2.*CK+1.)*U20
0237      704 A(N,J)=DEXP(U2K)
0238      DO 705 N=I3P,I2
0239      J=N+I6
0240      K=N-I3P
0241      CK=K
0242      U1K=(2.*CK+1.)*U10
0243      705 A(N,J)=DEXP(U1K)
0244      DO 706 N=I3P,I2

```

```

0245      J=N+I3
0246      A(N,J)=-BC1
0247      DO 707 N=I2P,I23
0248      J=N+I3
0249      707 A(N,J)=-BC2
0250      DO 708 N=I2P,I23
0251      J=N-I6
0252      K=N-I2P
0253      CK=K
0254      U2K=(2.*CK+1.)*U20
0255      708 A(N,J)=DEXP(U2K)
0256      JM=1
0257      DO 709 N=I23P,I56M
0258      J=N-I23
0259      JP=J+1
0260      K=N-I23P
0261      L=K+1
0262      M=K-1
0263      CK=K
0264      CL=L
0265      CM=M
0266      U1K=(CK+.5)*U10
0267      U1M=(CM+.5)*U10
0268      U1L=(CL+.5)*U10
0269      T11=DSINH(U10)-(2.*CK+1.)*DCOSH(U10)
0270      A(N,JM)=CK*GAME1*DEXP(-U1M)
0271      A(N,J)=GAME1*T11*DEXP(-U1K)
0272      A(N,JP)=CL*GAME1*DEXP(-U1L)
0273      709 JM=J
0274      DO 710 N=I23P,I56M
0275      J=N-I2
0276      JM=J-1
0277      JP=J+1
0278      K=N-I23P
0279      L=K+1
0280      M=K-1
0281      CL=L
0282      CK=K
0283      CM=M
0284      U1K=(CK+.5)*U10
0285      U1M=(CM+.5)*U10
0286      U1L=(CL+.5)*U10
0287      T12=DSINH(U10)+(2.*CK+1.)*DCOSH(U10)
0288      A(N,J)=GAME1*T12*DEXP(U1K)
0289      A(N,JM)=-GAME1*CK*DEXP(U1M)
0290      710 A(N,JP)=-GAME1*CL*DEXP(U1L)
0291      DO 711 N=I23P,I56M

```



```

0292      J=N-13
0293      JM=J-1
0294      JP=J+1
0295      K=N-I23P
0296      L=K+1
0297      M=K-1
0298      CL=L
0299      CK=K
0300      CM=M
0301      U1K=(CK+.5)*U10
0302      U1M=(CM+.5)*U10
0303      U1L=(CL+.5)*U10
0304      T11=DSINH(U10)-(2.*CK+1.)*DCOSH(U10)
0305      A(N,J)=T11*DEXP(-U1K)
0306      A(N,JM)=CK*DEXP(-U1M)
0307      A(N,JP)=CL*DEXP(-U1L)
0308      DO 712 N=I23P,I56M
0309      J=N-I6
0310      JM=J-1
0311      JP=J+1
0312      K=N-I23P
0313      L=K+1
0314      M=K-1
0315      CL=L
0316      CM=M
0317      CK=K
0318      U1K=(CK+.5)*U10
0319      U1M=(CM+.5)*U10
0320      U1L=(CL+.5)*U10
0321      T12=DSINH(U10)+(2.*CK+1.)*DCOSH(U10)
0322      A(N,J)=T12*DEXP(U1K)
0323      A(N,JM)=-CK*DEXP(U1M)
0324      A(N,JP)=-CL*DEXP(U1L)
0325      DO 713 N=I23P,I56M

0326      J=N
0327      JM=J-1
0328      JP=J+1
0329      K=N-I23P
0330      L=K+1
0331      M=K-1
0332      CK=K
0333      CM=M
0334      CL=L
0335      U1K=(CK+.5)*U10
0336      U1M=(CM+.5)*U10
0337      U1L=(CL+.5)*U10
0338      T11=DSINH(U10)-(2.*CK+1.)*DCOSH(U10)

```

711

712

```

0339      A(N,J)=-GAM11*T11*DEXP(-U1K)
0340      A(N,JM)=-GAM11*CK*DEXP(-U1M)
0341 713    A(N,JP)=-GAM11*CL*DEXP(-U1L)
0342      JM=1
0343      DO 714 N=156,IM2
0344      J=N-156M
0345      JP=J+1
0346      K=N-156
0347      L=K+1
0348      M=K-1
0349      CL=L
0350      CK=K
0351      CM=M
0352      U2L=(CL+.5)*U20
0353      U2K=(CK+.5)*U20
0354      U2M=(CM+.5)*U20
0355      T21=-DSINH(U20)-(2.*CK+1.)*DCOSH(U20)
0356      A(N,JM)=CK*GAME2*DEXP(U2M)
0357      A(N,J)=GAME2*T21*DEXP(U2K)
0358      A(N,JP)=CL*GAME2*DEXP(U2L)
0359 714    JM=J
0360      DO 715 N=156,IM2
0361      J=N-123M
0362      JM=J-1
0363      JP=J+1
0364      K=N-156
0365      L=K+1
0366      M=K-1
0367      CK=K
0368      CL=L
0369      CM=M
0370      U2L=(CL+.5)*U20
0371      U2K=(CK+.5)*U20
0372      U2M=(CM+.5)*U20
0373      T22=-DSINH(U20)+(2.*CK+1.)*DCOSH(U20)
0374      A(N,J)=+GAME2*T22*DEXP(-U2K)
0375      A(N,JM)=-CK*GAME2*DEXP(-U2M)
0376 715    A(N,JP)=-CL*GAME2*DEXP(-U2L)
0377      DO 718 N=156,IM2
0378      J=N+1
0379      JM=J-1
0380      JP=J+1
0381      K=N-156
0382      L=K+1
0383      M=K-1

```

```

0384      CL=L
0385      CM=M
0386      CK=K
0387      U2L=(CL+.5)*U20
0388      U2K=(CK+.5)*U20
0389      U2M=(CM+.5)*U20
0390      T22=-DSINH(U20)+(2.*CK+1.)*DCOSH(U20)
0391      A(N,J)=-GAMI2*T22*DEXP(-U2K)
0392      A(N,JM)=GAMI2*CK*DEXP(-U2M)
0393      718 A(N,JP)=GAMI2*CL*DEXP(-U2L)
0394      N=I-1
0395      DO 719 J=I6P,I3
0396      K=J+I3
0397      A(N,J)=GAME1
0398      719 A(N,K)=1.0
0399      N=I
0400      DO 720 J=1,I6
0401      K=J+I3
0402      A(N,J)=GAME2
0403      720 A(N,K)=1.
0404      DO 722 J=I3P,I2
0405      K=J+I6
0406      B(J)=CP1
0407      722 B(K)=CP2
0408      DO 721 J=1,I3
0409      K=J+I23
0410      B(J)=-TS
0411      721 B(K)=0.0
0412      IF (IR-1)7000,7000,7001
0413      7001 CONTINUE
C      VOLATILE BOUNDARY CONDITION
0414      DO 7725 J=1,I6
0415      K=J+I6
0416      B(J)=-TS
0417      7725 B(K)=-TS
0418      DO 7716 N=I56,IM2
0419      J=N-I2M
0420      K=N-I56
0421      JM=J-1
0422      JP=J+1
0423      L=K+1
0424      M=K-1
0425      CL=L
0426      CK=K
0427      CM=M
0428      U2L=(CL+.5)*U20
0429      U2K=(CK+.5)*U20

```

```

0430      U2M=(CM+.5)*U20
0431      T21=-DSINH(U20)-(2.*CK+1.)*DCOSH(U20)
0432      A(N,J)=T21*DEXP(U2K)
0433      A(N,JM)=CK*DEXP(U2M)
0434      7716  A(N,JP)=CL*DEXP(U2L)
0435      DO 7717 N=I56,I2M
0436      J=N-I3M
0437      K=N-I56
0438      JM=J-1
0439      JP=J+1
0440      L=K+1
0441      M=K-1
0442      CL=L
0443      CM=M
0444      CK=K
0445      U2L=(CL+.5)*U20
0446      U2K=(CK+.5)*U20
0447      U2M=(CM+.5)*U20
0448      T22=-DSINH(U20)+(2.*CK+1.)*DCOSH(U20)
0449      A(N,J)=T22*DEXP(-U2K)
0450      A(N,JM)=-CK*DEXP(-U2M)
0451      7717  A(N,JP)=-CL*DEXP(-U2L)
0452      GO TO 7002
0453      7000  CONTINUE
C      INERT BOUNDARY CONDITION
0454      DO 716 N=I2P,I23M
0455      J=N-I6
0456      K=N-I2P
0457      JM=J-1
0458      JP=J+1
0459      L=K+1
0460      M=K-1
0461      CL=L
0462      CK=K
0463      CM=M
0464      U2L=(CL+.5)*U20
0465      U2K=(CK+.5)*U20
0466      U2M=(CM+.5)*U20
0467      T21=-DSINH(U20)-(2.*CK+1.)*DCOSH(U20)
0468      A(N,J)=T21*DEXP(U2K)
0469      A(N,JM)=CK*DEXP(U2M)
0470      716   A(N,JP)=CL*DEXP(U2L)
0471      DO 717 N=I2P,I23M
0472      J=N
0473      JM=J-1
0474      JP=J+1
0475      K=N-I2P

```

```

0476      L=K+1
0477      M=K-1
0478      CL=L
0479      CM=M
0480      CK=K
0481      U2L=(CL+.5)*U20
0482      U2K=(CK+.5)*U20
0483      U2M=(CM+.5)*U20
0484      T22=-DSINH(U20)+(2.*CK+1.)*DCOSH(U20)
0485      A(N,J)=T22*DEXP(-U2K)
0486      A(N,JM)=-CK*DEXP(-U2M)
0487 717    A(N,JP)=-CL*DEXP(-U2L)
0488      A(I23,I23)=0
0489      DO 723 J=I3P,I2
0490      N=I23
0491 723    A(N,J)=1.
0492      DO 724 N=I2P,I23
0493      J=N+I3
0494      A(N,J)=0.
0495 724    B(N)=0.
0496      DO 725 J=1,I6
0497      N=I
0498      K=J+I3
0499      A(N,J)=0.0
0500 725    A(N,J)=1.
0501 7002  CONTINUE
0502      ETS=.1E-16
0503      CALL DGELG(B,A,I,1,ETS,IER)
0504      WRITE(3,13)
0505      WRITE(3,1)(B(N),N=1,I)
0506      WRITE(3,2)
0507      WRITE(3,25)
0508      WRITE(3,5)ETS,IER
0509      WRITE(3,2)
C    CALCULATION OF PHORETIC FORCE
0510      RHO1=0.
0511      RHO2=0.
0512      DO 403 N=1,10
0513      CN=N-1
0514      NC=I3+N
0515      ND=I2+N
0516      XN=CN+.5
0517      UN1=XN*U10
0518      UN2=XN*U20
0519      PHE1(N)=B(NC)*DEXP(-UN1)+B(ND)*DEXP(UN1)
0520      PHE2(N)=B(NC)*DEXP(UN2)+B(ND)*DEXP(-UN2)
0521      RHO1=RHO1+PHE1(N)
0522 403    RHO2=RHO2+PHE2(N)

```

```

0523 RHO1=DSQRT(DCOSH(U10)-1.)*RHO1+RHR
0524 RHO2=DSQRT(DCOSH(U20)-1.)*RHO2+RHR
0525 WRITE(3,37)
0526 WRITE(3,1)RHO1,RHO2
0527 WRITE(3,2)
0528 DO 401 N=1,7
0529 CN=N
0530 NC=I3+N
0531 ND=I2+N
0532 XN=CN+.5
0533 XM=CN-.5
0534 XK=CN+1.5
0535 TXM=2.*XM
0536 TXN=2.*XN
0537 TXK=2.*XK
0538 UN1=XN*U10
0539 UN2=XN*U20
0540 UM1=XM*U10
0541 UM2=XM*U20
0542 UK1=XK*U10
0543 UK2=XK*U20
0544 NP=N+2
0545 NPP=N+1
0546 PHI1(N)=PHE1(N)+PHE1(NP)-2.*DCOSH(U10)*PHE1(NPP)
0547 PHI2(N)=PHE2(N)+PHE2(NP)-DCOSH(U20)*PHE2(NPP)*2.
0548 SIM=DSINH(UM1+UM2)
0549 GAP1(N)=+(1.-SLP)*AD*(AM/RHG)*PHI1(N)*N*(N+1)*DCOSH(UM1)*SIM/TXN
0550 GAP2(N)=GAP1(N)*PHI2(N)*DCOSH(UM2)/DCOSH(UM1)/PHI1(N)
0551 GAP1(N)=GAP1(N)/(DCOSH(UM1)*SIM)
0552 GAP2(N)=GAP2(N)/(DCOSH(UM2)*SIM)
0553 AA(1,1)=DCOSH(UM1)
0554 AA(1,2)=DSINH(UM1)
0555 AA(1,3)=DCOSH(UK1)
0556 AA(1,4)=DSINH(UK1)
0557 AA(3,1)=DCOSH(UM2)
0558 AA(3,2)=-DSINH(UM2)
0559 AA(3,3)=DCOSH(UK2)
0560 AA(3,4)=-DSINH(UK2)
0561 AA(2,1)=TXM*DSINH(UM1)
0562 AA(2,2)=TXM*DCOSH(UM1)
0563 AA(2,3)=TXK*DSINH(UK1)
0564 AA(2,4)=TXK*DCOSH(UK1)
0565 AA(4,1)=-TXM*DSINH(UM2)
0566 AA(4,2)=TXM*DCOSH(UM2)
0567 AA(4,3)=-TXK*DSINH(UK2)
0568 AA(4,4)=TXK*DCOSH(UK2)

```

```

0569      BB(1)=0.
0570      BB(3)=0.
0571      BB(2)=GAP1(N)
0572      BB(4)=GAP2(N)
0573      CALL DGELG(BB,AA,4,1,ETS,IER)
0574      WRITE(3,25)
0575      WRITE(3,1)ETS,IER
0576      H(N) = BB(1)
0577      E(N) = BB(2)
0578      C(N) = BB(3)
0579      D(N) = BB(4)
0580      F1=-TXN*(H(N)+E(N)+C(N)+D(N))+F1
0581      F2=TXN*(H(N)-E(N)+C(N)-D(N))+F2
401      WRITE(3,2)
0583      COF=-VIS*3.1416*2.*SQ2/AM
0584      F1=COF*F1
0585      F2=COF*F2
0586      WRITE(3,39)
0587      WRITE(3,1)(H(N),E(N),C(N),D(N),N=1,7)
0588      WRITE(3,2)
0589      WRITE(3,52)
0590      WRITE(3,1)F1,F2
0591      FOR2(JE)=F2*.1000D+09
0592      FOR1(JE)=F1*.1000D+09
0593      V1=F1*10000./(6.*PI*VIS*A1)
0594      V2=F2*10000./(6.*PI*VIS*A2)
0595      WRITE(3,54)
0596      WRITE(3,1)V1,V2
0597      WRITE(3,2)
C      BOUNDARY CONDITION CHECK
      DO 666 N=1,7
0598      CN=N
0599      UM1=(CN-.5)*U10
0600      UM2=(CN-.5)*U20
0601      UK1=(CN+1.5)*U10
0602      UK2=(CN+1.5)*U20
0603      SI1=H(N)*DCOSH(UM1)+E(N)*DSINH(UM1)+C(N)*DCOSH(UK1)+D(N)
0604      1*DSINH(UK1)
0605      SI2=H(N)*DCOSH(UM2)-E(N)*DSINH(UM2)+C(N)*DCOSH(UK2)-D(N)
      1*DSINH(UK2)
0606      SIM=DSINH(UM1+UM2)
0607      DSI1=(2.*CN-1.)*(H(N)*DSINH(UM1)+E(N)*DCOSH(UM1))+(2.*CN+3.)*
      1(C(N)*DSINH(UK1)+D(N)*DCOSH(UK1))-GAP1(N)
0608      DSI2=(2.*CN-1.)*(-H(N)*DSINH(UM2)+E(N)*DCOSH(UM2))+(2.*CN+3.)

```

```

0609      1*(-C(N)*DSINH(UK2)+D(N)*DCOSH(UK2))-GAP2(N)
0610      WRITE(3,53)
0611      WRITE(3,1)SI1,SI2,DSI1,DSI2
0612      WRITE(3,2)
666 0613      CONTINUE
799 0614      CONTINUE
0615      FG1=-4./3.*PI*A1*A1*A1*GRAV
0616      FG2=-4./3.*PI*A2*A2*A2*GRAV
0617      FPL2=FG2*DCOS(ALF)+F2
0618      FPL1=FG1*DCOS(ALF)+F1
0619      FPD1=FG1*DSIN(ALF)
0620      FPD2=FG2*DSIN(ALF)
C      CALCULATION OF STOKES RESISTANCE
0621      COF=-VIS*3.1416*2.*SQ2/AM
0622      S1=6.*3.1416*VIS*A1
0623      S2=S1*A2/A1
0624      R1=FPL1/S1
0625      R2=FPL2/S2
0626      COF1=COF/S1
0627      COF2=COF/S2
0628      Y1P=0.
0629      Y2P=0.
0630      Y1PP=0.
0631      Y2PP=0.
0632      DO 402 N=1,10
0633      XM=N-.5
0634      XK=N+1.5
0635      TXM=2.*XM
0636      TXK=2.*XK
0637      XP=N+.5
0638      TXP=2.*XP
0639      UM1=XM*U10
0640      UM2=XM*U20
0641      UK1=XK*U10
0642      UK2=XK*U20
0643      UP1=XP*U10
0644      UP2=XP*U20
0645      DELT=4.*DSINH(UP1+UP2)*DSINH(UP1+UP2)-TXP*TXP*DSINH(U10+U20)
1*DSINH(U10+U20)
DELT=1./DELT

```



```

0646      GK=AM*AM*N*(N+1)/(SQ2*TXM*TXP*TXK)
0647      AS(N)=DEL*GK*TXK*(4.*DEXP(-UP1-UP2)*DSINH(UP1+UP2)+TXP*TXP*
1DEXP(U10+U20)*DSINH(U10+U20)+2.*TXM*DSINH(UP1+UP2)*DCOSH(UP1-UP2)
2-2.*TXP*DSINH(UK1+UK2)*DCOSH(UM1-UM2)-TXP*TXM*DSINH(U10+U20)*
3DCOSH(U10-U20))
0648      BS(N)=-DEL*TXK*GK*(2.*TXM*DSINH(UP1+UP2)*DSINH(UP1-UP2)-2.*TXP*
1DSINH(UK1+UK2)*DSINH(UM1-UM2)+TXP*TXM*DSINH(U10+U20)
2*DSINH(U10-U20))
0649      CS(N)=-DEL*TXM*GK*(4.*DEXP(-UP1-UP2)*DSINH(UP1+UP2)-TXP*TXP
1*DEXP(-U10-U20)*DSINH(U10+U20)+2.*TXP*DSINH(UM1+UM2)
1*DCOSH(UK1-UK2)-2.*TXK*DSINH(UP1+UP2)*DCOSH(UP1-UP2)+TXP*TXK*
2DSINH(U10+U20)*DCOSH(U10-U20))
0650      DS(N)=DEL*TXM*GK*(2.*TXP*DSINH(UM1+UM2)*DSINH(UK1-UK2)-2.*TXK
1*DSINH(UP1+UP2)*DSINH(UP1-UP2)+TXP*TXK*DSINH(U10+U20)
2*DSINH(U10-U20))
0651      ASP(N)=TXK*GK*(2.*TXM*DSINH(UP1+UP2)*DSINH(UP1-UP2)-2.*TXP*
1DSINH(UK1+UK2)*DSINH(UM1-UM2)-TXM*TXP*DSINH(U10+U20)*
2DSINH(U10-U20))*DEL
0652      BSP(N)=-TXK*GK*(-4.*DEXP(-UP1-UP2)*DSINH(UP1+UP2)-TXP*TXP*
1DEXP(U10+U20)*DSINH(U10+U20)+2.*TXM*DSINH(UP1+UP2)*DCOSH(UP1-UP2)
2-2.*TXP*DSINH(UK1+UK2)*DCOSH(UM1-UM2)+TXM*TXP*DSINH(U10+U20)
3*DCOSH(U10-U20))*DEL
C      CSP IS ARTIFICALLY SUPRESSED TO ZERO
0653      CSP(N)=-TXM*GK*(2.*TXP*DSINH(UM1+UM2)*DSINH(UK1-UK2)-2.*TXK
1*DSINH(UP1+UP2)*DSINH(UP1-UP2)-TXP*TXK*DSINH(U10+U20)*
2DSINH(U10-U20))*DEL
0654      DSP(N)=TXM*GK*(-4.*DEXP(-UP1-UP2)*DSINH(UP1+UP2)+TXP*TXP*
1DEXP(-U10-U20)*DSINH(U10+U20)+2.*TXP*DSINH(UM1+UM2)*DCOSH(UK1-UK2)
2-2.*TXK*DSINH(UP1+UP2)*DCOSH(UP1-UP2)-TXP*TXK*DSINH(U10+U20)
3*DCOSH(U10-U20))*DEL
0655      SJ1(N)=AS(N)*DCOSH(UM1)+BS(N)*DSINH(UM1)+CS(N)*DCOSH(UK1)
1+DS(N)*DSINH(UK1)
0656      SJ2(N)=AS(N)*DCOSH(UM2)-BS(N)*DSINH(UM2)+CS(N)*DCOSH(UK2)
1-DS(N)*DSINH(UK2)
0657      SJ3(N)=TXM*(AS(N)*DSINH(UM1)+BS(N)*DCOSH(UM1))+TXK*(CS(N)*
1DSINH(UK1)+DS(N)*DCOSH(UK1))
0658      SJ4(N)=TXM*(-AS(N)*DSINH(UM2)+BS(N)*DCOSH(UM2))+TXK*(-CS(N)*
1DSINH(UK2)+DS(N)*DCOSH(UK2))
0659      DM1(N)=ASP(N)*DCOSH(UM1)+BSP(N)*DSINH(UM1)+CSP(N)*DCOSH(UK1)+
1DSP(N)*DSINH(UK1)

```

```

0660      DM2(N)=ASP(N)*DCOSH(UM2)-BSP(N)*DSINH(UM2)+CSP(N)*DCOSH(UK2)-
0661      1DSP(N)*DSINH(UK2)
0662      DM3(N)=TXM*(ASP(N)*DSINH(UM1)+BSP(N)*DCOSH(UM1))+TXK*(CSP(N)*
0663      1DSINH(UK1)+DSP(N)*DCOSH(UK1))
0664      DM4(N)=TXM*(-ASP(N)*DSINH(UM2)+BSP(N)*DCOSH(UM2))+TXK*(-CSP(N)*
0665      1DSINH(UK2)+DSP(N)*DCOSH(UK2))
0666      RE1(N)=GK*(TXK*DEXP(-UM1)-TXM*DEXP(-UK1))
0667      RE2(N)=GK*(TXK*DEXP(-UM2)-TXM*DEXP(-UK2))
0668      RE3(N)=-GK*TXK*TXM*(DEXP(-UM1)-DEXP(-UK1))
0669      RE4(N)=GK*TXK*TXM*(DEXP(-UM2)-DEXP(-UK2))
0670      Y1P=COF1*TXP*(AS(N)+BS(N)+CS(N)+DS(N))+Y1P
0671      Y2P=Y2P+COF2*TXP*(AS(N)-BS(N)+CS(N)-DS(N))
0672      Y1PP=Y1PP+COF1*TXP*(ASP(N)+BSP(N)+CSP(N)+DSP(N))
0673      402 Y2PP=Y2PP+COF2*TXP*(ASP(N)-BSP(N)+CSP(N)-DSP(N))
0674      DUN=Y1PP*Y2P-Y2PP*Y1P
0675      VP=(R2*Y1PP-R1*Y2PP)/DUN
0676      VPP=(+R1*Y2P-R2*Y1P)/DUN
0677      V1PL=VP-VPP
0678      V2PL=VP+VPP
0679      WRITE(3,47)
0680      WRITE(3,3)Y1P,Y2P,Y1PP,Y2PP,VP,VPP
0681      WRITE(3,2)
0682      WRITE(3,42)
0683      WRITE(3,1)(AS(N),BS(N),CS(N),DS(N),N=1,10)
0684      WRITE(3,2)
0685      WRITE(3,46)
0686      WRITE(3,1)(ASP(N),BSP(N),CSP(N),DSP(N),N=1,10)
0687      WRITE(3,41)
0688      YAMB=R1/V1PL
0689      WRITE(3,1)DA,V1PL,V2PL,YAMB
0690      WRITE(3,2)
0691      DO 992 N=1,10
0692      BSJ1=SJ1(N)+RE1(N)
0693      BSJ2=SJ2(N)+RE2(N)
0694      BSJ3=SJ3(N)+RE3(N)
0695      BSJ4=SJ4(N)+RE4(N)

```

```

0693      BDM1=DM1(N)-RE1(N)
0694      BDM2=DM2(N)+RE2(N)
0695      BDM3=DM3(N)-RE3(N)
0696      BDM4=DM4(N)+RE4(N)
0697      WRITE(3,66)
0698      WRITE(3,1)BSJ1,BSJ2,BSJ3,BSJ4
0699      WRITE(3,67)
0700      992 WRITE(3,1)BDM1,BDM2,BDM3,BDM4
C MOTION PERPENDICULAR TO LINE OF CENTERS
0701      L=JF-2
0702      L8=L/8
0703      L4=L/4
0704      L2=L/2
0705      L23=2*L/3
0706      L56=5*L/6
0707      L512=5*L/12
0708      L58=5*L/8
0709      L78=7*L/8
0710      L34=3*L/4
0711      L38=3*L/8
0712      L712=7*L/12
0713      L1112=11*L/12
0714      LP=L+1
0715      LL=L+2
0716      L8M=L8-1
0717      L2P=L2+1
0718      L8P=L8+1
0719      L4P=L4+1
0720      L512P=L512+1
0721      L58P=L58+1
0722      FPD1P=FPD1/(6*PI*VIS*A1)
0723      FPD2P=FPD2/(6*PI*VIS*A2)
0724      DO 299 NR=1,LL
0725      DO 299 NC=1,LL
0726      AP(NR,NC)=0.
0727      299 BP(NR)=0.
0728      DO 298 N=1,L4
0729      UN1=(N-.5)*U10
0730      UN2=(N-.5)*U20
0731      AN1(N)=DEXP(UN1)/DSINH(U10)
0732      AN2(N)=-DEXP(-UN2)/DSINH(U20)
0733      BN1(N)=DEXP(-UN1)/DSINH(U10)
0734      298 BN2(N)=-DEXP(UN2)/DSINH(U20)

```

```

0735      NCM=3
0736      NRM=3
0737      DO 300 NR=1,L8
0738      NC=NR+L8
0739      NRP=NR+1
0740      NRQ=NRP+1
0741      AP(NR,NR)=-2.*DCOSH(U10)*AN1(NRP)
0742      AP(NR,NC)=-2.*DCOSH(U10)*BN1(NRP)
0743      AP(NC,NC)=-2.*DCOSH(U20)*BN2(NRP)
0744      AP(NC,NR)=-2.*DCOSH(U20)*AN2(NRP)
0745      COD=2.*(NR-1)/(2*NR-1.)
0746      AP(NR,NRM)=COD*AN1(NR)
0747      AP(NR,NCM)=COD*BN1(NR)
0748      AP(NC,NRM)=COD*AN2(NR)
0749      AP(NC,NCM)=COD*BN2(NR)

0750      NCM=NC
0751      NRM=NR
0752      NCP=NC+1
0753      CODE=2.*(NR+2)/(2*NR+3.)
0754      AP(NR,NRP)=CODE*AN1(NRQ)
0755      AP(NR,NCP)=CODE*BN1(NRQ)
0756      AP(NC,NRP)=CODE*AN2(NRQ)
0757      AP(NC,NCP)=CODE*BN2(NRQ)
0758      NCC=NC+L8
0759      NCD=NC+L4
0760      UN1=(NR+.5)*U10
0761      UN2=(NR+.5)*U20
0762      AP(NR,NCC)=-DEXP(UN1)
0763      AP(NR,NCD)=-DEXP(-UN1)
0764      AP(NC,NCC)=-DEXP(-UN2)
0765      AP(NC,NCD)=-DEXP(UN2)
300      WRITE(3,2)
0767      DO 301 NR=1,L8P
0768      UN1=(NR-.5)*U10
0769      UN2=(NR-.5)*U20
0770      NC=NR+L8
0771      NRE=NR+L4
0772      NCE=NR+L2
0773      NCF=NR+L2+L8P
0774      AP(NRE,NCE)=-DEXP(UN1)
0775      AP(NRE,NCF)=-DEXP(-UN1)
0776      NP=NR+1
0777      NR1=L4+NR
0778      NREE=NR1+L8P
0779      AP(NREE,NCE)=-DEXP(-UN2)
0780      AP(NREE,NCF)=-DEXP(UN2)
0781      CAD=NR*(NR+1.)/(2*NR+1.)

```

```

0782      AP(NR1,NR)=CAD*AN1(NP)
0783      AP(NR1,NC)=CAD*BN1(NP)
0784      NR2=NR1+L8P
0785      AP(NR2,NR)=CAD*AN2(NP)
0786      AP(NR2,NC)=CAD*BN2(NP)
0787      AP(NR1,LP)=2.*SQ2*DEXP(-UN1)
0788      301 AP(NR2,LL)=2.*SQ2*DEXP(-UN2)
0789      DO 302 NR=2,L8
0790      UN1=(NR+.5)*U10
0791      UN2=(NR+.5)*U20
0792      NR1=L4P+NR
0793      NC=NR-1
0794      NC1=NC+L8
0795      NR2=NR1+L8P
0796      NRP=NR+1
0797      NCP=NRP+L8
0798      CAD=-NR*(NR-1)/(2.*NR-1)
0799      AP(NR1,NC)=CAD*AN1(NR)
0800      AP(NR1,NC1)=CAD*BN1(NR)
0801      AP(NR2,NC)=CAD*AN2(NR)
0802      AP(NR2,NC1)=CAD*BN2(NR)
0803      NR3=NR+L2P
0804      NR4=NC+L4+L8P+L8P+L8M
0805      NC2=NC+L2+L8P+L8P
0806      NC3=NR+L78
0807      AP(NR3,NC2)=-DEXP(UN1)
0808      AP(NR3,NC3)=-DEXP(-UN1)
0809      AP(NR4,NC2)=-DEXP(-UN2)
0810      AP(NR4,NC3)=-DEXP(UN2)
0811      CID=1./(2*NR-1.)
0812      AP(NR3,NC)=CID*AN1(NR)
0813      AP(NR3,NC1)=CID*BN1(NR)
0814      AP(NR4,NC)=CID*AN2(NR)
0815      AP(NR4,NC1)=CID*BN2(NR)
0816      NPP=NRP+1
0817      CODE=-1./(2*NR+3.)
0818      AP(NR3,NRP)=CODE*AN1(NPP)
0819      AP(NR3,NCP)=CODE*BN1(NPP)
0820      AP(NR4,NRP)=CODE*AN2(NPP)
0821      302 AP(NR4,NCP)=CODE*BN2(NPP)
0822      DO 303 NR=1,L8
0823      NR2=NR+L78
0824      NR1=NR+L34
0825      NC1=NR+L8
0826      NC2=NR+L4
0827      NC3=NR+L38

```

```

0828      NC4=NR+L2
0829      NC5=NR+L2+L8P
0830      NC4P=NC4+1
0831      NC5P=NC5+1
0832      CID=2.*(2*NR+1.)
0833      AP(NR1, NR)=CID
0834      AP(NR2, NC1)=-CID
0835      AP(NR1, NC2)=5.
0836      AP(NR2, NC3)=5.
0837      AP(NR1, NC4)=-1.
0838      AP(NR2, NC5)=-1.
0839      AP(NR1, NC4P)=2.
0840      AP(NR2, NC5P)=2.
303      DO 304 NR=1, L8M
0841      NR1=NR+L34
0842      NR2=NR+L78
0843      NP=NR+1
0844      NR3=NP+L34
0845      NR4=NP+L78
0846      NC1=NR+L4
0847      NC2=NR+L38
0848      CN=NR
0849      AP(NR3, NC1)=-CN
0850      AP(NR4, NC2)=-CN
0851      NC3=NP+L4
0852      NC4=NP+L38
0853      AP(NR1, NC3)=NP+1.
0854      AP(NR2, NC4)=NP+1.
0855      NC5=L2+NR+2
0856      NC6=L2+L8P+2+NR
0857      AP(NR1, NC5)=-1.
0858      AP(NR2, NC6)=-1.
0859      NC7=L2+L8P+L8P+NR
0860      NC8=NC7+L8M
0861      AP(NR1, NC7)=(NR+2)*(NR+3.)
0862      AP(NR2, NC8)=(NR+2)*(NR+3.)
0863      AP(NR3, NC7)=-2.*NR*(NP+2)
0864      AP(NR4, NC8)=-2.*NR*(NP+2)
0865
0866      NR5=NR+L8
0867      AP(NR3, NR)=-2.*NR
0868      AP(NR4, NR5)=+2.*NR
0869      NP1=NP+L8
0870      AP(NR1, NP)=-2.*(NR+2)
0871      AP(NR2, NP1)=+2.*(NR+2)
304      DO 305 NR=3, L8
0872      NR1=NR+L34
0873

```

```

0874      NM=NR-2
0875      NR2=NR+L78
0876      NC1=NM+L2+L8P+L8P
0877      NC2=NC1+L8M
0878      AP(NR1,NC1)=(NR-2.)*(NR-1)
0879      305 AP(NR2,NC2)=(NR-2.)*(NR-1)
0880      FACT1=+6.*PI*VIS*A1*2.*SQ2/3.*DSINH(U10)
0881      FACT2=+6.*PI*VIS*A2*2.*SQ2/3.*DSINH(U20)
0882      DO 306 NC=1,L8P
0883      NC1=NC+L2
0884      NC2=NC1+L8P
0885      AP(LP,NC1)=FACT1
0886      306 AP(LL,NC2)=FACT2
0887      BP(LP)=FPD1
0888      BP(LL)=FPD2
0889      AP(L8,L8P)=0
0890      AP(L8,L4P)=0
0891      AP(L4,L4P)=0
0892      AP(L4,L8P)=0
0893      AP(L58P,L8P)=0
0894      AP(L58P,L4P)=0
0895      AP(L34,L8P)=0
0896      AP(L34,L4P)=0
0897      L513=L4+L8P
0898      L713=L4+L8P+L8P
0899      AP(L513,L8P)=0
0900      AP(L513,L4P)=0
0901      AP(L713,L8P)=0
0902      AP(L713,L4P)=0
0903      ETS=.1E-16
0904      CALL DGELG(BP,AP,LL,1,ETS,IER)
0905      WRITE(3,55)
0906      WRITE(3,3)(BP(N),N=1,LL)
0907      WRITE(3,2)
0908      WRITE(3,25)
0909      WRITE(3,5)ETS,IER
0910      WRITE(3,2)
0911      WRITE(3,71)
0912      WRITE(3,1)FPD1,FPD2,FPD1P,FPD2P
0913      CHK1=C1*BP(97)+C5*BP(98)
0914      CHK2=C2*BP(97)+C6*BP(98)
0915      WRITE(3,72)
0916      WRITE(3,1)CHK1,CHK2
0917      DO 2000 N=1,L8
0918      UN1=(N+.5)*U10
0919      UN2=(N+.5)*U20
0920      NLA=N
0921      NLB=N+L8
0922      NLC=N+L4
0923      NLD=N+L38

```

```

0924      AB1(N)=(BP(NLA)*DEXP(UN1)+BP(NLB)*DEXP(-UN1))/DSINH(U10)
0925      AB2(N)=-(BP(NLA)*DEXP(-UN2)+BP(NLB)*DEXP(UN2))/DSINH(U20)
0926      ARAP(N)=BP(NLA)*DEXP(UN1)/DSINH(U10)
0927      ARBP(N)=BP(NLB)*DEXP(-UN1)/DSINH(U10)
0928      ARAM(N)=-BP(NLA)*DEXP(-UN2)/DSINH(U20)
0929      ARBM(N)=-BP(NLB)*DEXP(UN2)/DSINH(U20)
0930      CD1(N)=(BP(NLC)*DEXP(UN1)+BP(NLD)*DEXP(-UN1))
0931      2000 CD2(N)=+(BP(NLC)*DEXP(-UN2)+BP(NLD)*DEXP(UN2))
0932      DO 2001 N=1,L8P
0933      UN1=(N-.5)*U10
0934      UN2=(N-.5)*U20
0935      NLE=N+L2
0936      NLF=N+L2+L8P
0937      EF1(N)=(BP(NLE)*DEXP(UN1)+BP(NLF)*DEXP(-UN1))
0938      2001 EF2(N)=+(BP(NLE)*DEXP(-UN2)+BP(NLF)*DEXP(UN2))
0939      DO 2002 N=1,L8M
0940      UN1=(N+1.5)*U10
0941      UN2=(N+1.5)*U20
0942      NLG=N+L2+L8P+L8P
0943      NLH=N+L2+L8P+L8P+L8M
0944      GH1(N)=(BP(NLG)*DEXP(UN1)+BP(NLH)*DEXP(-UN1))
0945      2002 GH2(N)=+(BP(NLG)*DEXP(-UN2)+BP(NLH)*DEXP(UN2))
0946      DO 2003 N=1,L8
0947      AB1(L8P)=0
0948      AB2(L8P)=0
0949      K=N+1
0950      M=N-1
0951      IF (M)2004,2004,2005
0952      2004 M=N
0953      2005 ARB1=-2*(N+2)/(2*N+3.)*AB1(K)
0954      BRB1=-2*(N+2)/(2*N+3.)*AB2(K)
0955      ARB2=2*DCOSH(U10)*AB1(N)
0956      BRB2=2*DCOSH(U20)*AB2(N)
0957      ARB3=-(N-1)/(2*N-1.)*AB1(M)*2.
0958      BRB3=-(N-1)/(2*N-1.)*AB2(M)*2.
0959      CND11=CD1(N)+ARB1+ARB2+ARB3
0960      CND12=CD2(N)+BRB1+BRB2+BRB3
0961      WRITE(3,78) N
0962      WRITE(3,82)
0963      2003 WRITE(3,1)CND11,CND12
0964      DO 2006 N=1,L8P
0965      UN1=(N-.5)*U10
0966      UN2=(N-.5)*U20
0967      K=N+1
0968      M=N-1
0969      MM=M-1

```



```

0970      AB1(L8P)=C
0971      AB2(L8P)=0
0972      ARB1=-2*BP(LP)*SQ2*DEXP(-UN1)
0973      BRB1=-2*BP(LL)*SQ2*DEXP(-UN2)
0974      ARB2=-(M+1)*(M+2)/(2*M+3.)*AB1(N)
0975      BRB2=-(M+1)*(M+2)/(2*M+3.)*AB2(N)
0976      IF (MM)2007,2007,2008
0977      2007 MM=N
0978      2008 ARB3=M*(M-1)/(2*M-1.)*AB1(MM)
0979      BRB3=M*(M-1)/(2*M-1.)*AB2(MM)
0980      CND21=EF1(N)+ARB1+ARB2+ARB3
0981      CND22=EF2(N)+BRB1+BRB2+BRB3
0982      WRITE(3,78) N
0983      WRITE(3,83)
0984      2006 WRITE(3,1)CND21,CND22
0985      DO 2009 N=1,L8M
0986      K=N+1
0987      KK=N+2
0988      AB1(L8P)=C
0989      AB2(L8P)=0
0990      ARAP(L8P)=0.
0991      ARAM(L8P)=0.
0992      ARBM(L8P)=0.
0993      ARBP(L8P)=0.
0994      ARB1=GH1(N)
0995      BRB1=GH2(N)
0996      CID=1./(2.*K-1.)
0997      CODE=1./(2.*K+3.)
0998      ARB2A=-CID*ARAP(N)
0999      BRB2A=-CID*ARAM(N)
1000      ARB2B=-CID*ARBP(N)
1001      BRB2B=-CID*ARBM(N)
1002      ARB2=ARB2A+ARB2B
1003      BRB2=BRB2A+BRB2B
1004      ARB3A=CODE*ARAP(KK)
1005      BRB3A=CODE*ARAM(KK)
1006      ARB3B=CODE*ARBP(KK)
1007      BRB3B=CODE*ARBM(KK)
1008      ARB3=ARB3A+ARB3B
1009      BRB3=BRB3A+BRB3B
1010      CND31=ARB1+ARB2+ARB3
1011      CND32=BRB1+BRB2+BRB3
1012      WRITE(3,78) N
1013      WRITE(3,84)
1014      2009 WRITE(3,1)CND31,CND32

```

```

1015      NUMB=L2P+L8P
1016      CND41=BP(L2P)
1017      CND42=-BP(NUMB)
1018      CND51=0.
1019      CND52=0.
1020      WRITE(3,85)
1021      DO 2010 N=1,L8
1022      M=N-1
1023      NLA=N
1024      NLB=N+L8
1025      NLE=N+L2
1026      NLF=N+L2+L8P
1027      NLEP=NLE+1
1028      NLFP=NLF+1
1029      CND41=CND41+2.*N*(N+1.)*BP(NLA)+(2.*N+1.)*BP(NLEP)
1030      CND42=CND42+2.*N*(N+1.)*BP(NLB)-(2.*N+1.)*BP(NLFP)
2010  WRITE(3,1)CND41,CND42
1032      DO 2011 N=1,L8P
1033      NLE=N+L2
1034      NLF=N+L2+L8P
1035      CND51=CND51+BP(NLE)
2011  CND52=CND52+BP(NLF)
1037      CND52=FPD2-6.*PI*VIS*A2*2.*SQ2/3.*DSINH(U20)*CND52
1038      CND51=FPD1-6.*PI*VIS*A1*2.*SQ2/3.*DSINH(U10)*CND51
1039      WRITE(3,86)
1040      WRITE(3,1)CND51,CND52
1041      V1PD=BP(LP)
1042      V2PD=BP(LL)
1043      V1X=V1PL*DSIN(ALF)-V1PD*DCOS(ALF)
1044      V1Z=V1PL*DCOS(ALF)+V1PD*DSIN(ALF)
1045      V2X=V2PL*DSIN(ALF)-V2PD*DCOS(ALF)
1046      V2Z=V2PL*DCOS(ALF)+V2PD*DSIN(ALF)
1047      VE1X(JE)=V1X
1048      VE2X(JE)=V2X
1049      VE1Z(JE)=V1Z
1050      VE2Z(JE)=V2Z
1051      WRITE(3,96)
1052      WRITE(3,1)V1X,V2X,V1Z,V2Z
1053      W1X(IND)=V1X
1054      IF (IND-2)602,603,603
1055      603  X1P=(W1X(1)+W1X(2))*DT/4.+W1X(1)*DT/2.+X1W(1)
1056      IF (DABS(X1P-X1W(2))-ERR1)604,604,605
1057      604  INDP=INDP+1
1058      IF (INDP-3)602,607,607
1059      607  IF (DABS((W1X(1)+W1X(3))/2.-W1X(2))-ERR2)608,608,602
1060      608  DT=2.*DT
1061      WRITE(3,98)

```

```

1062      IND=0
1063      GO TO 602
1064 605 DT=DT/2.
1065      WRITE(3,97)
1066      INDP=1
1067      X1=X1-V1XN*DT
1068      X2=X2-V2XN*DT
1069      Z1=Z1-V1ZN*DT
1070      Z2=Z2-V2ZN*DT
1071      GO TO 606
1072 602 CONTINUE
1073      X1=X1+V1X*DT
1074      X2=X2+V2X*DT
1075      Z1=Z1+V1Z*DT
1076      Z2=Z2+V2Z*DT
1077      DX=X1-X2
1078      DZ=Z1-Z2
1079      CHGX(JE)=DX
1080      CHGZ(JE)=DZ
1081      JE=JE+1
1082      IF (DZ) 833,834,833
1083 834 ALF=PI/2
1084      GO TO 836
1085 833 ALF=ATAN(DX/DZ)
1086 836 ARUG=DZ*DZ+DX*DX
1087      DA=DSQRT(ARUG)
1088      WRITE(3,63)
1089      WRITE(3,65)
1090      TT=TT+DT
1091      WRITE(3,1)DA,ALF,TT
1092      WRITE(3,95)
1093      WRITE(3,1)X1,X2,Z1,Z2
1094      WRITE(3,40)
1095      IF (DA-.0120)850,838,838
1096 850 GO TO 800
1097 838 CONTINUE
1098      JE=JE-1
1099      WRITE(3,99)
1100      WRITE(3,3)(DIST(N),ANGL(N),VE1X(N),VE1Z(N),VE2X(N),VE2Z(N),N=1,JE)
1101      WRITE(3,75)
1102      WRITE(3,100)(DIST(N),FOR1(N),FOR2(N),N=1,JE)
1103      WRITE(3,1)(CHGX(N),CHGZ(N),N=1,JE)
1104      RETURN
1105      END

```